

Global Simulations of Aerosol Size-Dependent In-Cloud Scavenging

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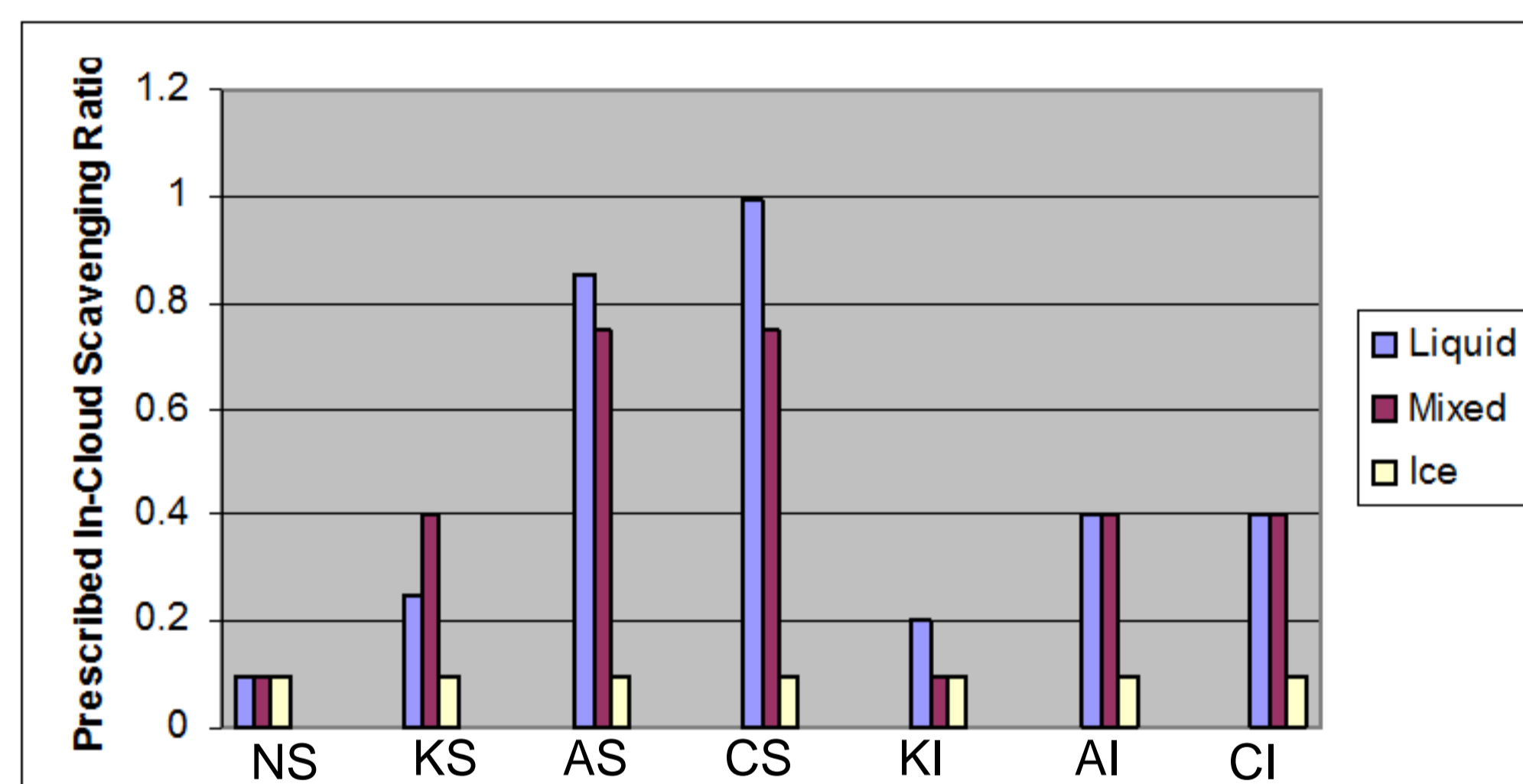
Introduction/Methodology:

Wet deposition processes strongly control aerosol three-dimensional distributions. Global models, such as GEOS-Chem typically treat wet deposition as a first order loss process, which depends on 1) a rate constant for the conversion of cloud water to precipitate, and 2) an aerosol cloud-borne fraction. In GEOS-Chem (Lui et al. (2001)) the soluble aerosol cloud-borne fraction is assumed to be unity for cloud droplets. In actuality, the cloud-borne fraction depends on a variety of parameters such as aerosol size/hygroscopicity, as well as cloud temperature/phase, and supersaturation. Alternatives to the assumption that 100% of in-cloud soluble aerosol is cloud-borne include:

- 1) prescribed cloud-borne fractions (e.g. Stier et al. (2005))
 - 2) diagnosed cloud-borne fractions based on certain parameters such as updraft velocity, supersaturation, aerosol size/hygroscopicity and cloud hydrometeor size (e.g. Croft et al. (2009))
 - 3) prognostic treatments of aerosol mass contained in cloud droplets and ice crystals (e.g. Hoose et al. (2008))
- This study compares predicted cloud-borne fractions and aerosol distributions for these three treatments of cloud-borne aerosol. All results presented are for 1-year simulations with the ECHAM5-HAM global climate model (Roeckner et al. (2003), Lohmann et al. (2007)).

Prescribed Cloud-Borne Fractions:

(As a function of cloud temperature and aerosol size)



ECHAM5-HAM log-normal modes:

Internally mixed modes:
 NS: Nucleation Soluble
 KS: Aitken Soluble
 AS: Accumulation Soluble
 CS: Coarse Soluble

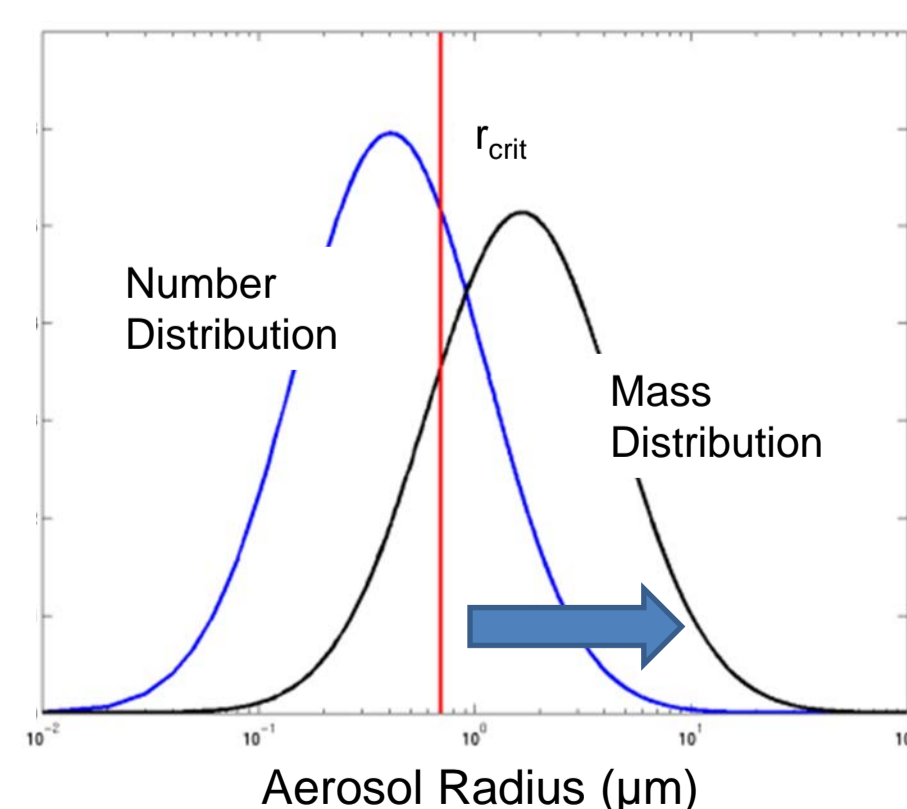
Externally mixed modes:
 KI: Aitken insoluble
 AI: Accumulation insoluble
 CI: Coarse insoluble

Mixed phase cloud: 238K < T < 273K

Size-Dependent Uptake into Cloud Hydrometeors:

Nucleation Scavenging:

Assume each cloud droplet and ice crystal scavenge one aerosol. Schematic below shows fractional scavenging of mass distribution is greater than for the number distribution, scavenge all mass for $r > r_{crit}$.

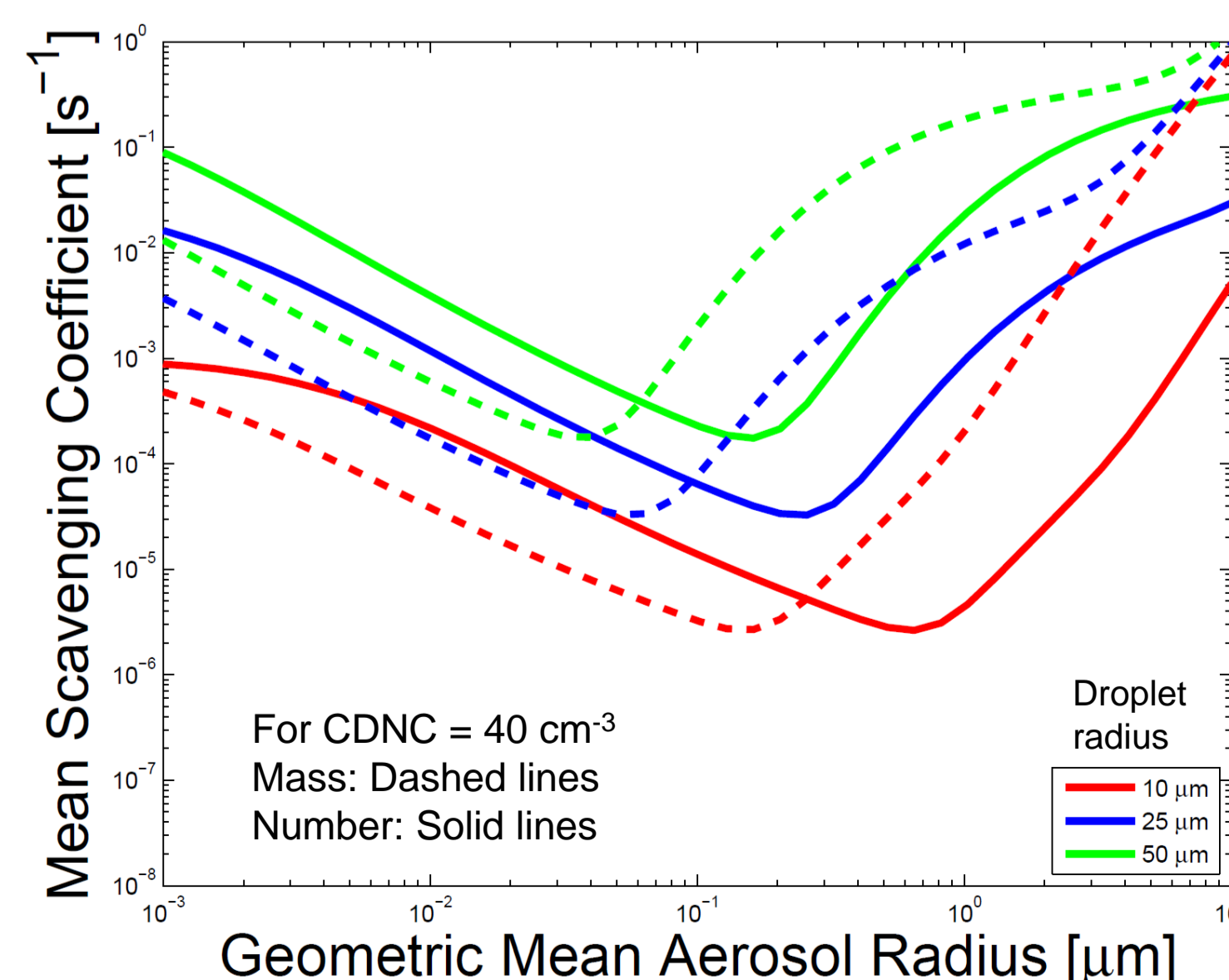


$$r_{crit} = r_g \left[\exp(\sqrt{2} \ln \sigma_g \cdot \text{erf}^{-1}(1 - (2 \cdot (CDNC + ICNC) \frac{frac_j}{N_{>35nm}})))) \right]$$

where $frac_j$ is the fraction of soluble mode j number having radius $> 35\text{nm}$
 $N_{>35nm}$ is the number of aerosols having radius $> 35\text{nm}$ for all soluble modes
 r_g and σ_g are parameters of the lognormal distribution, erf is the error function.

Impaction Scavenging:

Scavenging coefficients are found in look-up tables as a function of aerosol size, cloud droplet/ice crystal size and cloud droplet/ice crystal number concentration (CDNC/ICNC).

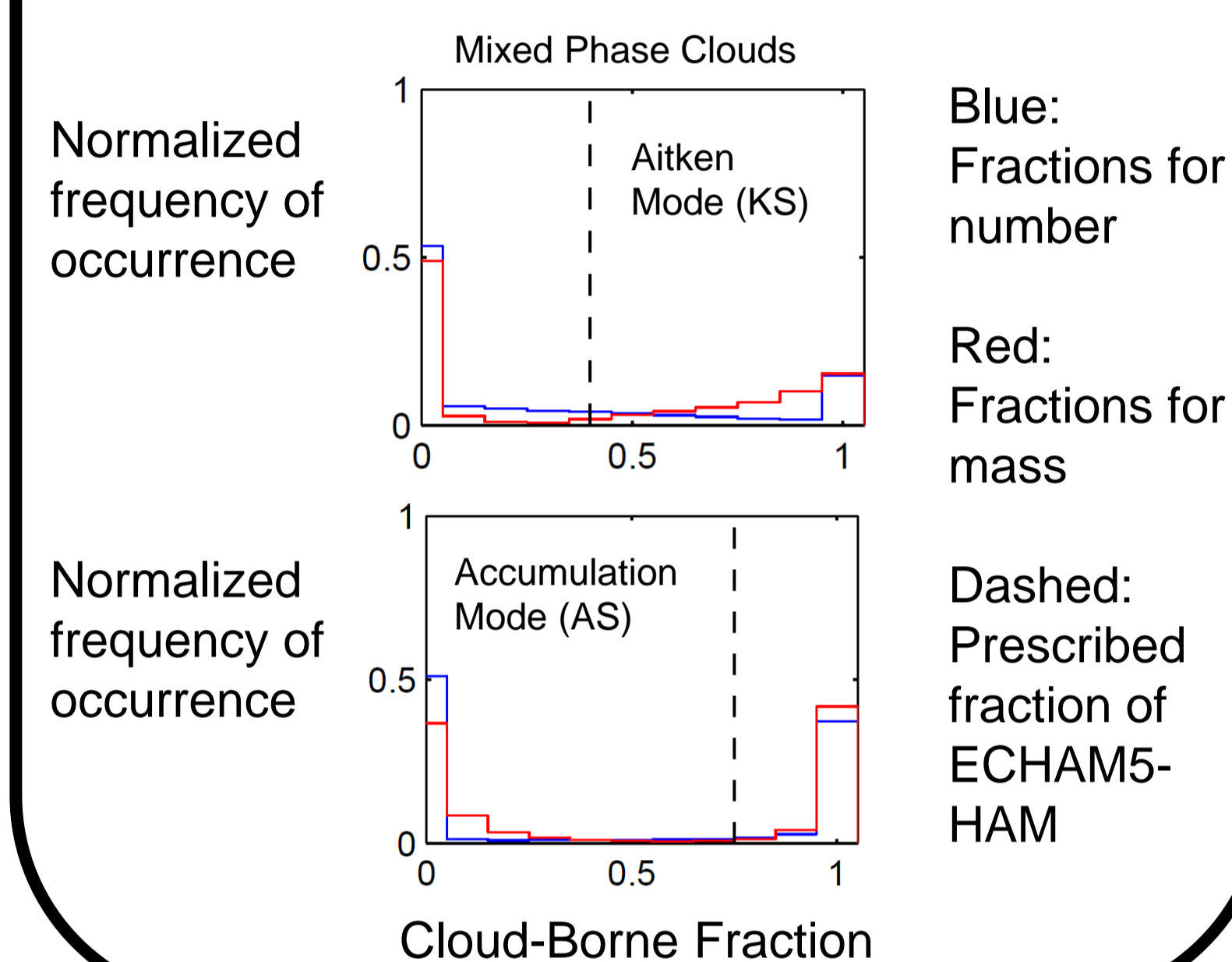


(Similar approach for aerosol-ice crystal collisions)

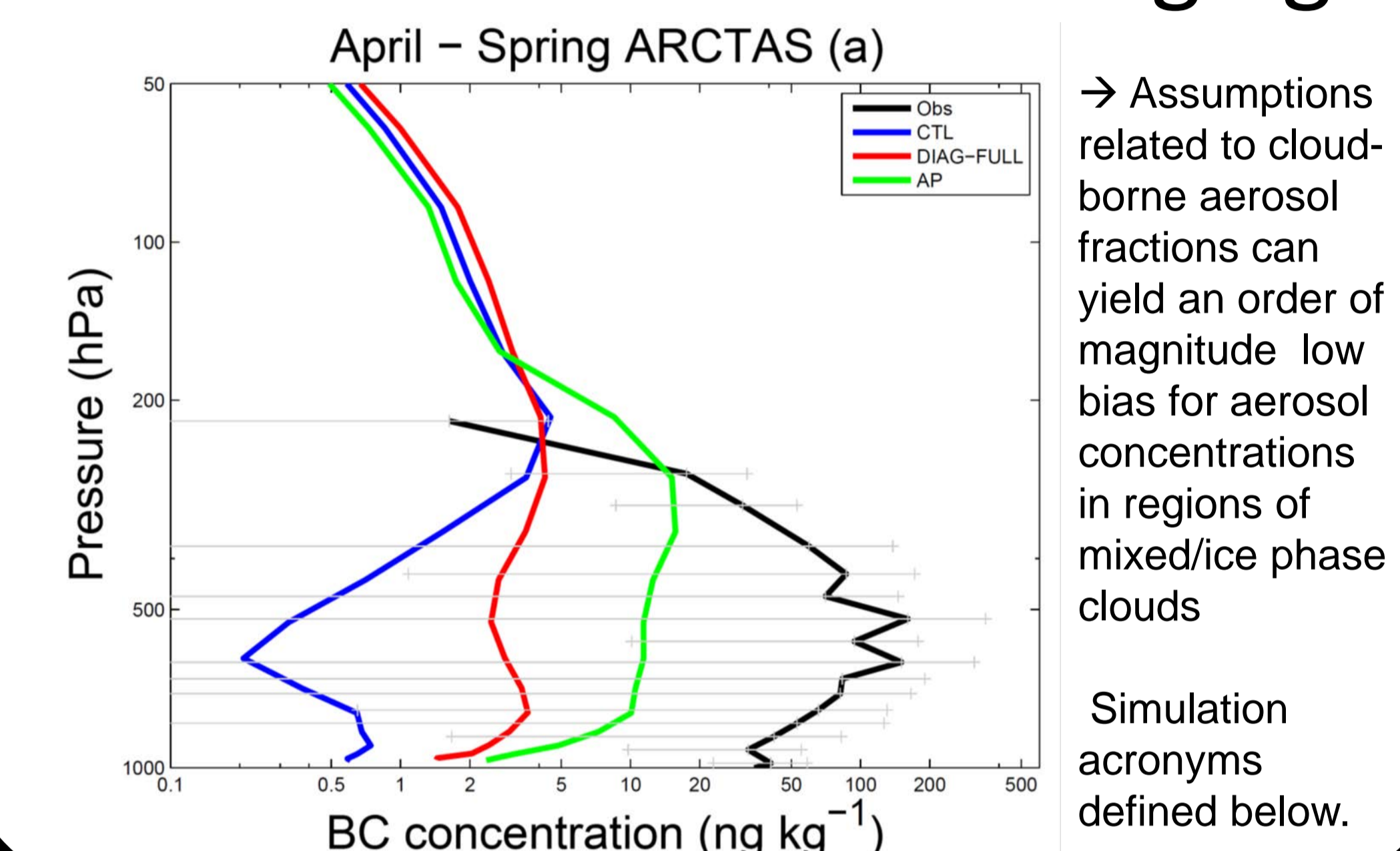
Summary:

- For fine mode soluble aerosols in mixed phase clouds, the cloud-borne fractions have wide variability, which prescribed fractions cannot capture. A mixed phase cloud-borne fraction of 0.8 for black carbon biased concentrations low by one order of magnitude in ECHAM5-HAM
- For coarse mode insoluble aerosols such as dust the parameterization of impaction scavenging is important for prediction of cloud-borne fractions
- Cloud-borne fractions should be assigned with attention to size/hygroscopicity, and cloud phase.

Diagnosed Cloud-Borne Fractions:



Sensitivity of Black Carbon Concentrations to Scavenging:



Marine Boundary Layer Size Distributions:

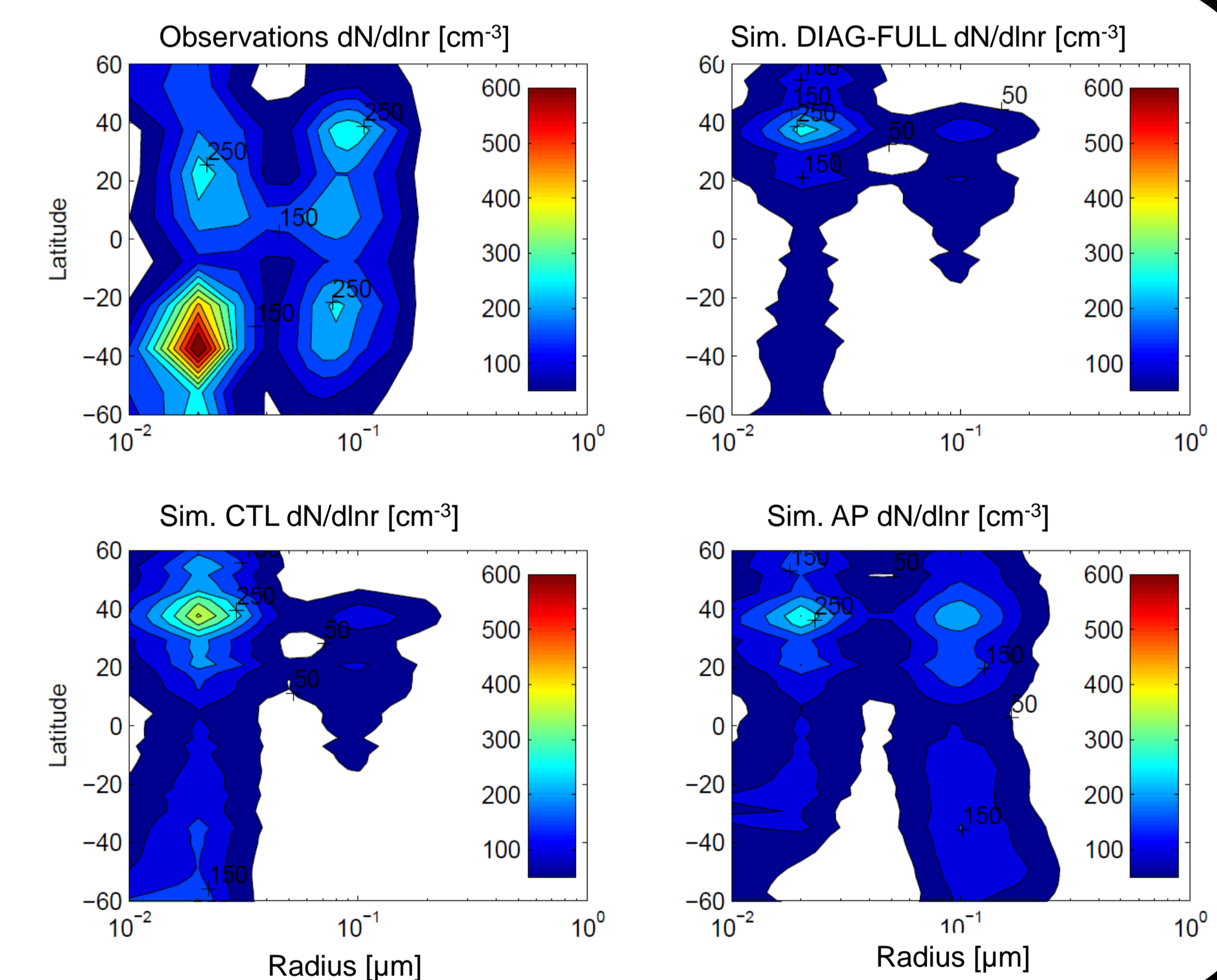
CTL: Standard model with prescribed cloud-borne aerosol fractions

DIAG-FULL: Diagnostic size-dependent uptake into cloud hydrometeors by nucleation and impaction scavenging

AP: In-droplet and in-crystal aerosol mass treated as prognostic species

Observations from Heintzenberg et al. (2000)

→ Prognostic scheme predicts accumulation mode in closest agreement with observations.



References:

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