

Sensitivity analysis of tropospheric ozone at middle and high latitudes to precursor emissions

Thomas Walker
University of Toronto
GEOS-Chem Users' Meeting
May 2, 2011



Outline

- Adjoint provides an efficient means of calculating sensitivities of the model state
- Examples where the adjoint model provides supplemental insights:
 - (1) Ozone transport to Arctic sites
 - (2) Surface ozone in eastern North America
 - (3) Midlatitude lightning NO_x emissions constraint

Given a small perturbation to the model inputs, the tangent linear model (TLM) can calculate the change in the model state

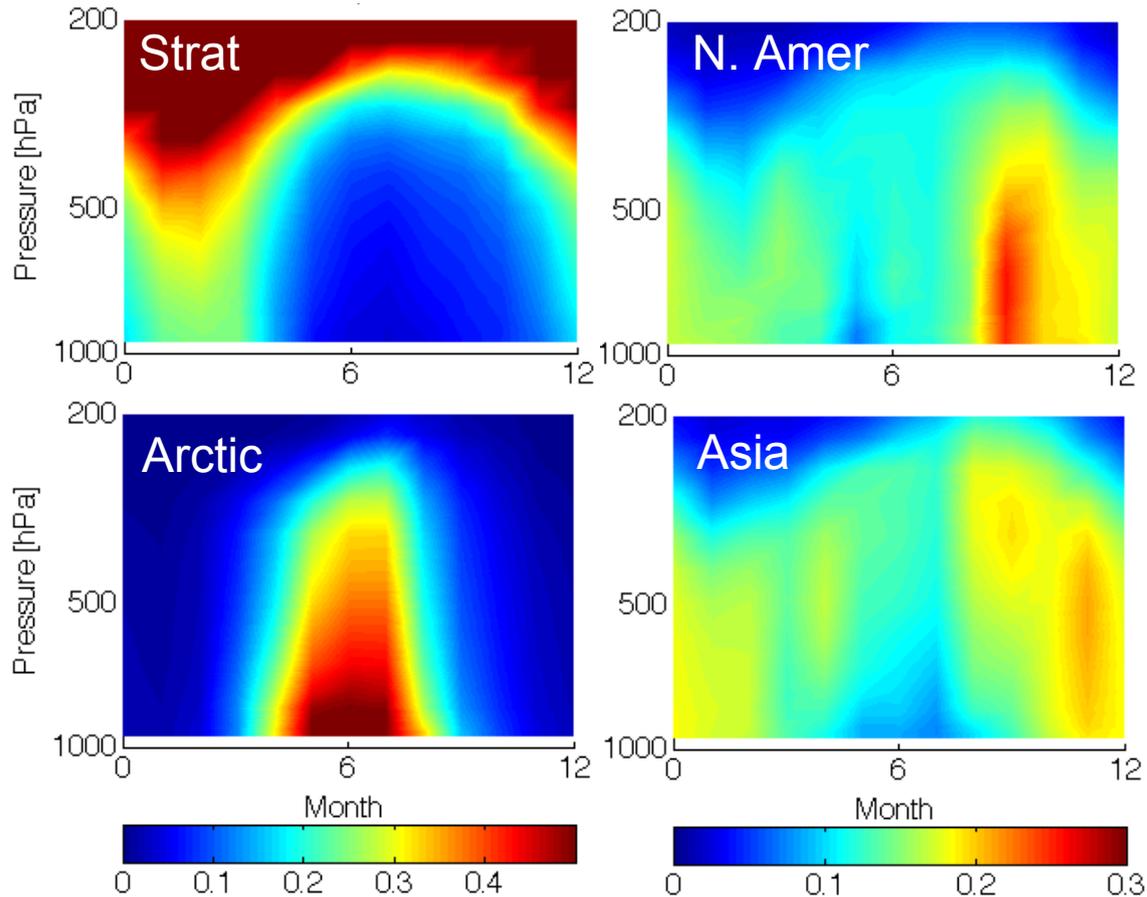
$$\Delta y = TLM(\Delta x)$$

Likewise, given a perturbed model state, the adjoint model (transpose of TLM) can calculate the required change in inputs

$$\Delta x = TLM^T(\Delta y)$$

Ozone Transport to Arctic Sites

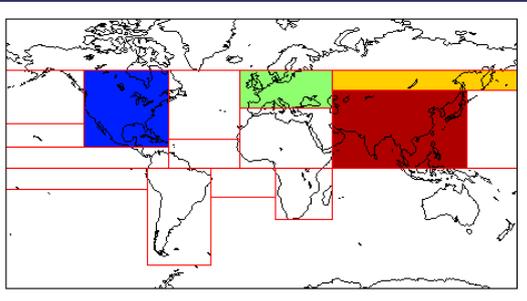
Fraction of total Ox above Eureka from



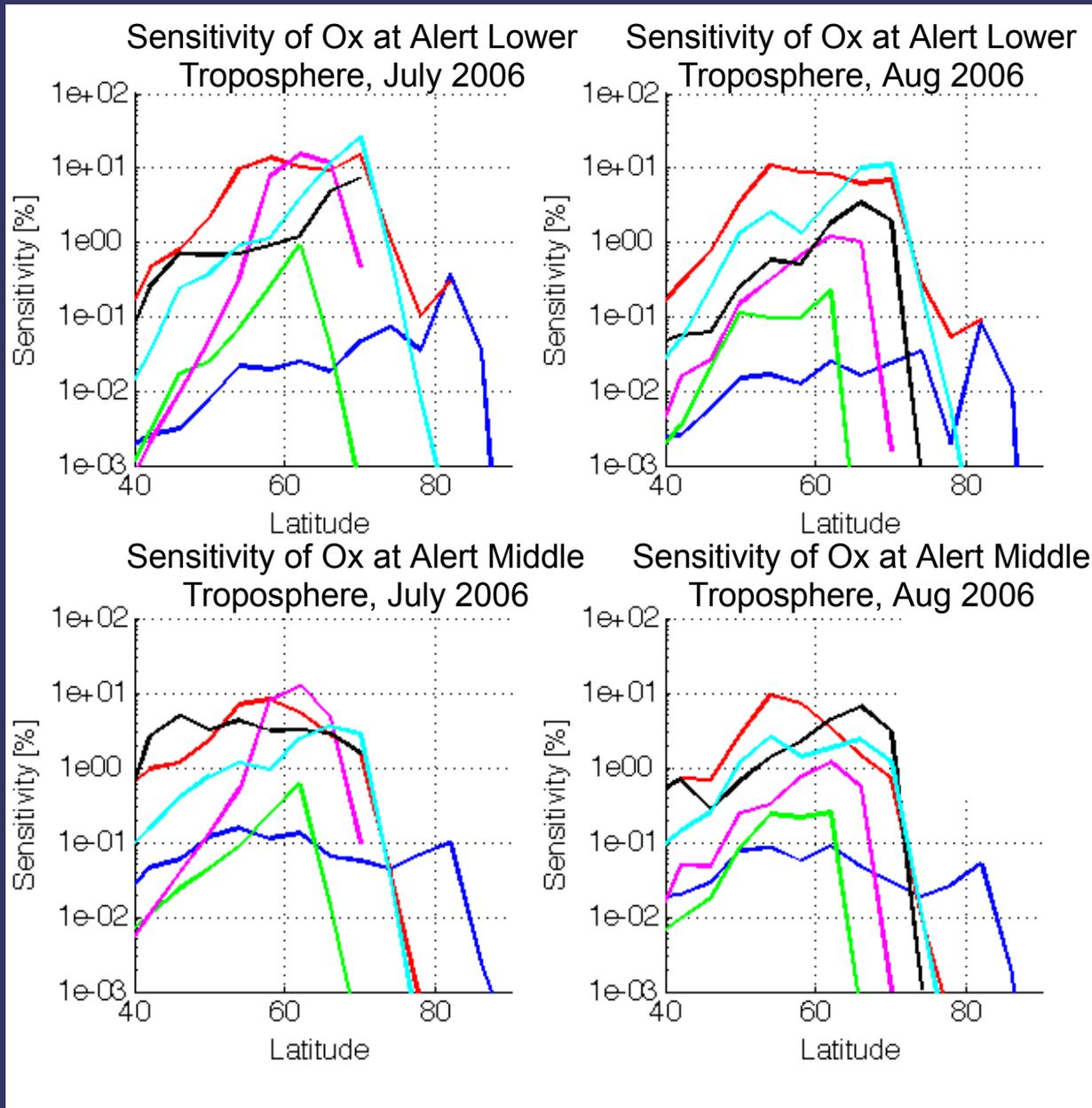
- One year of vertical profiles of tagged Ox tracers above Eureka, NU (80.0N, 86.2W)
- Stratospheric descent in spring; local influence inside “polar dome” in summer
- Transport from midlatitude source regions contribute 15-20% in spring and fall

... WHAT'S MISSING?

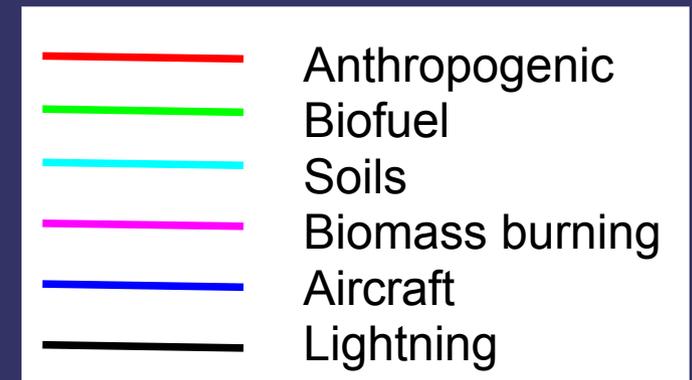
- No separation of precursor source types (how much from anthropogenic, biomass burning, etc...)
- “Traditional approach” - numerous forward simulations to isolate each source individually



Ozone Transport to Arctic Sites

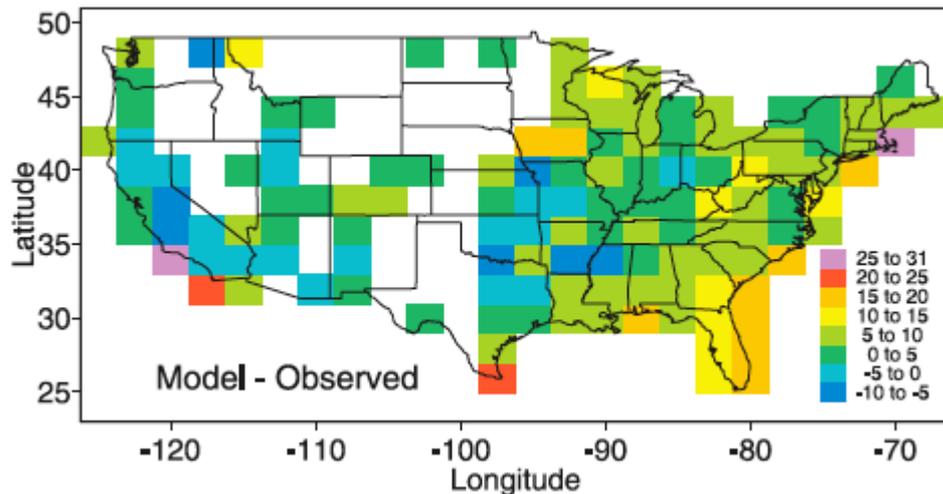


- Single adjoint model run calculates sensitivity of ozone at the Arctic site to different NO_x emissions sources
- Sensitivity is calculated for every grid box
- Strongest sensitivities are to anthropogenic (red), soils (cyan), lightning (black), and biomass burning (magenta)



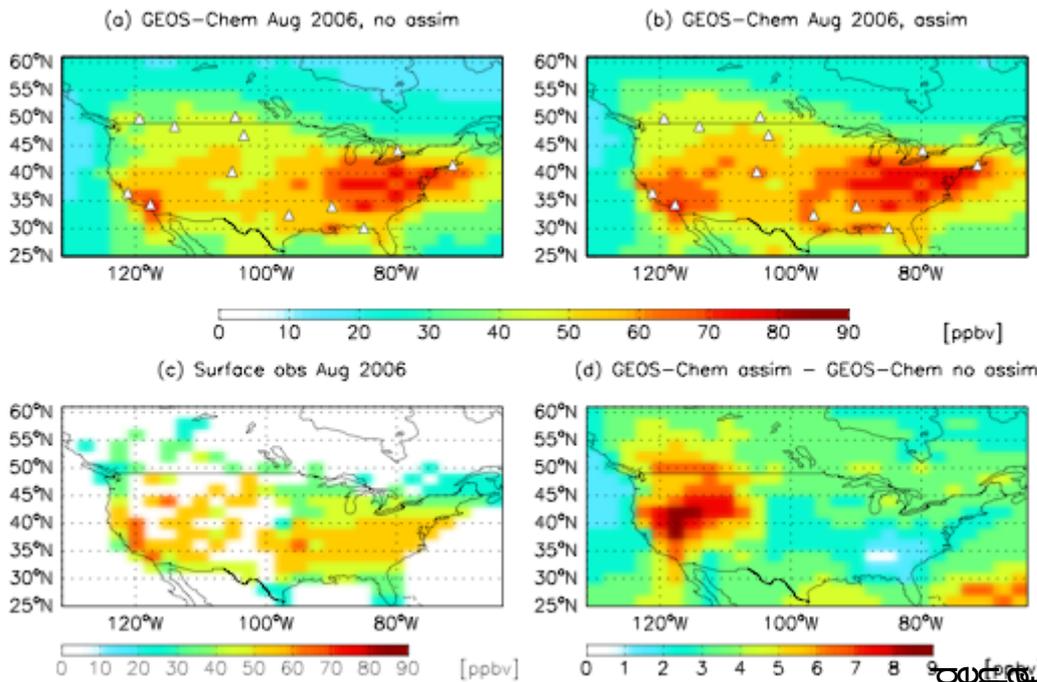
Watch for this in: Walker et al., (in prep.) *Impacts of midlatitude precursor emissions and local photochemistry on ozone abundances in the Arctic*

Surface Ozone in Eastern North America

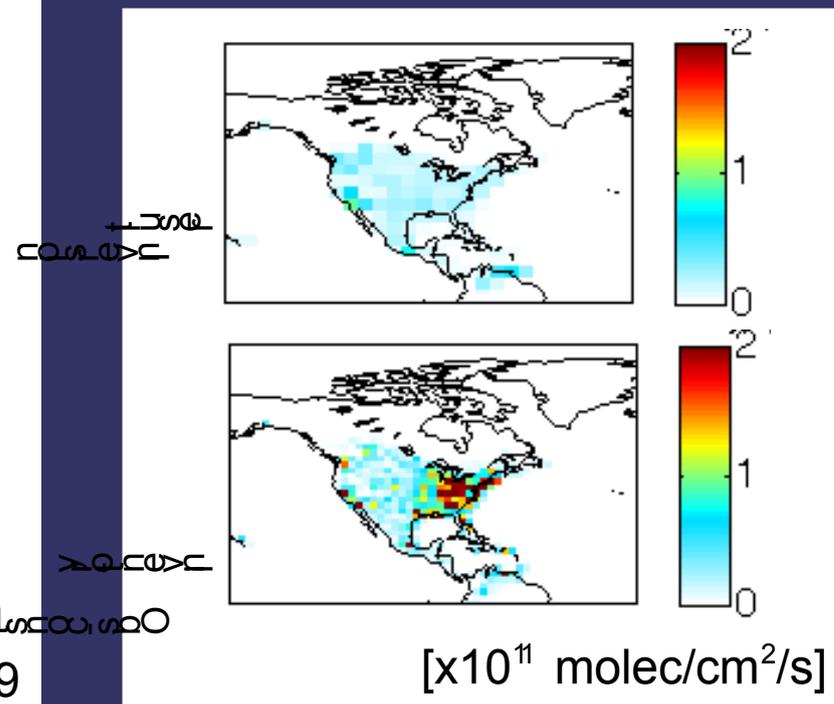


Fiore et al. 2002

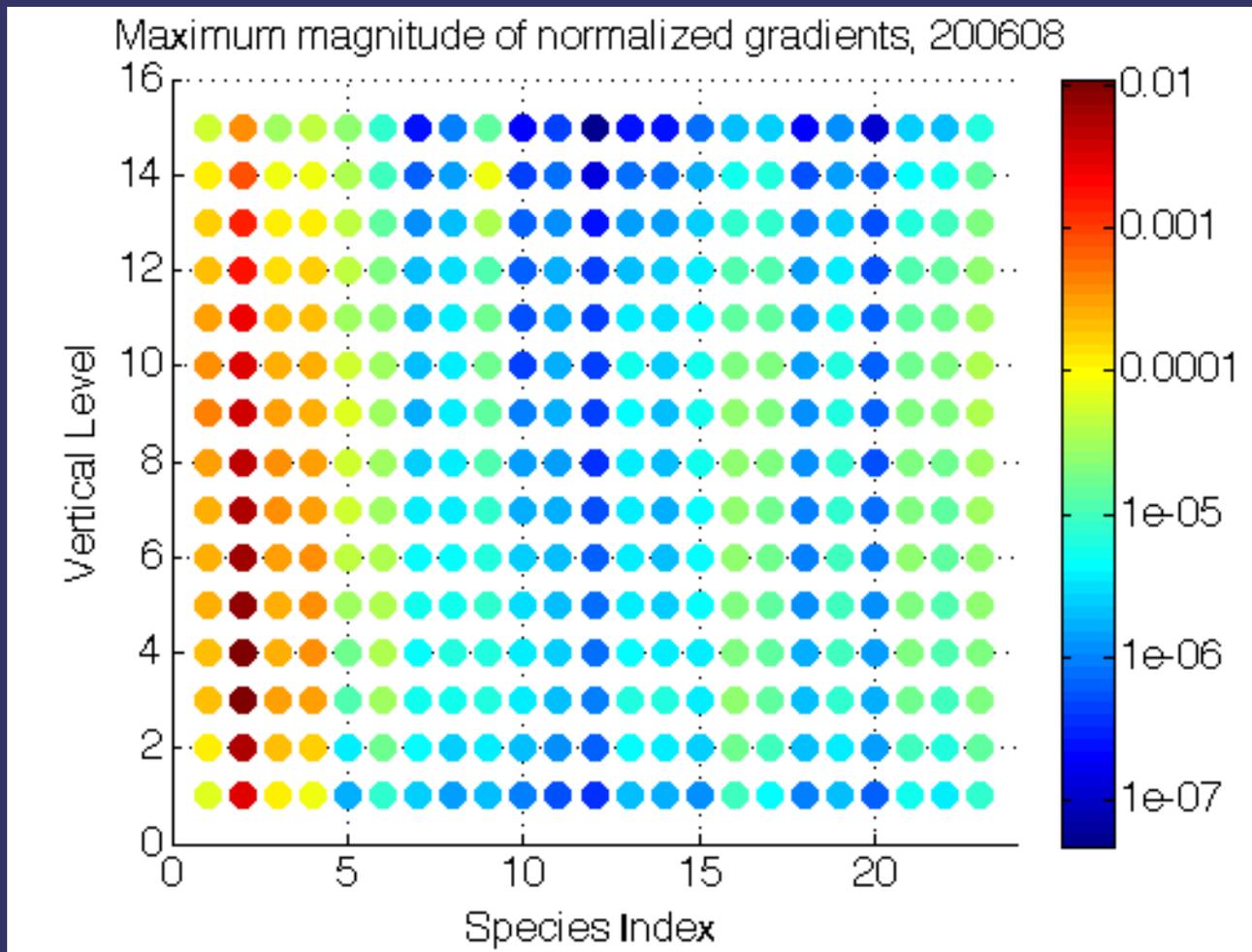
- GEOS-Chem has trouble reproducing surface ozone observations in some regions
- Constraining background tropospheric ozone with assimilated observations does not correct
- An inversion for NO_x emissions using surface ozone observations imposes an unrealistic constraint



Parrington et al. 2009



Surface Ozone in Eastern North America



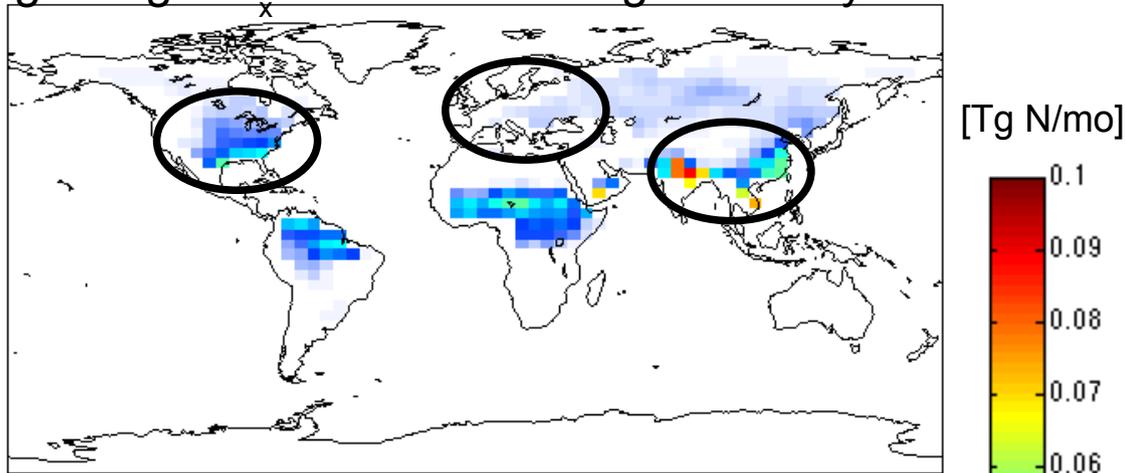
- | | |
|------------------------------------|------------------------------------|
| 1 – NO _x | 13 – MVK |
| 2 – Ox | 14 – MACR |
| 3 – PAN | 15 – PMN |
| 4 – CO | 16 – PPN |
| 5 – HNO ₃ | 17 – R ₄ N ₂ |
| 6 – ACET | 18 – PRPE |
| 7 – ALD ₂ | 19 – C ₃ H ₈ |
| 8 – ALK ₄ | 20 – CH ₂ O |
| 9 – ISOP | 21 – C ₂ H ₆ |
| 10 – H ₂ O ₂ | 22 – N ₂ O ₅ |
| 11 – MEK | 23 – HNO ₄ |
| 12 – RCHO | |

- Sensitivities to species concentrations can guide future studies (eg. Chai et al. (2006))
- Looking to develop sensitivities to more control parameters to offer more insights

Midlatitude Lightning NO_x Emissions Constraint

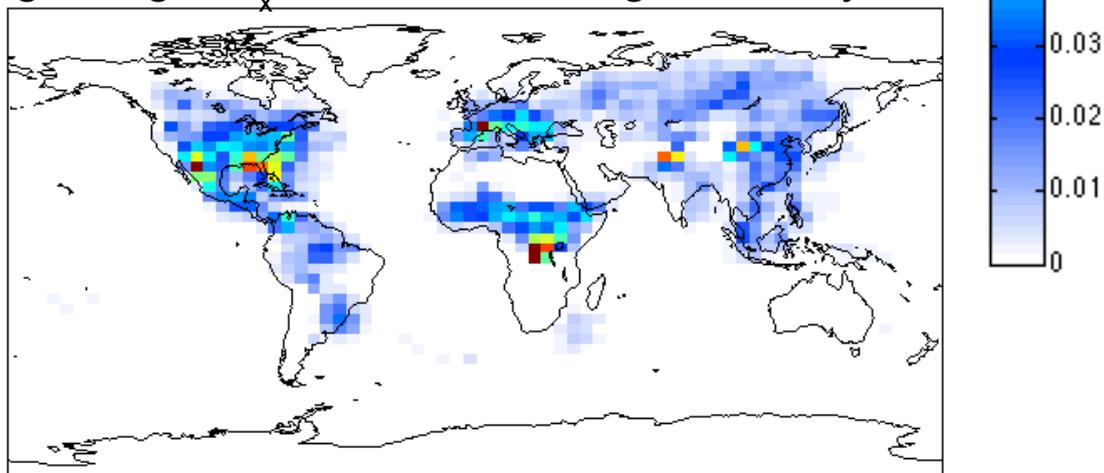
GEOS-Chem v7-02-04

Lightning NO_x source = 0.43 Tg N for July 2006



GEOS-Chem v8-01-04

Lightning NO_x source = 0.86 Tg N for July 2006

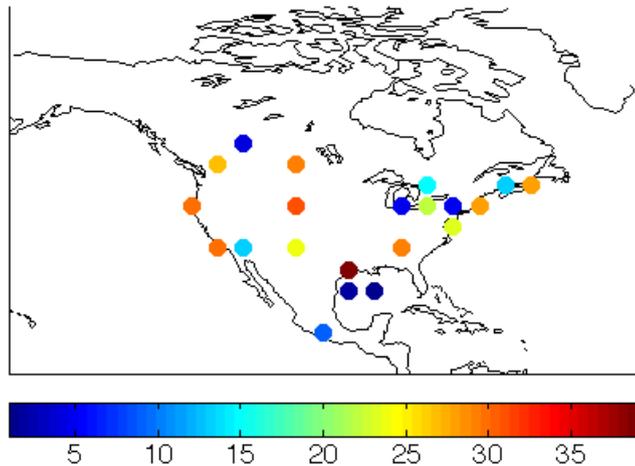


- Global lightning NO_x emissions in v7-02-04 and earlier follow Price & Rind (1992)
- Emissions in v8-01-04 and later are rescaled to match OTD-LIS climatology (Lee Murray)
- Rescaling produces a more realistic midlatitude ozone distribution

... CAN WE DO BETTER?

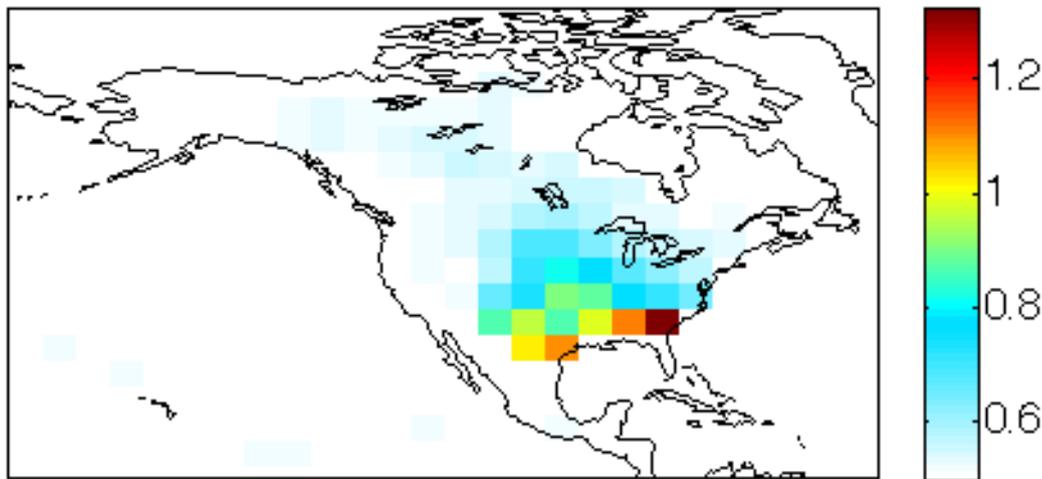
Midlatitude Lightning NO_x Emissions Constraint

Number of IONS06 ozonesondes in August 2006 (407 total)



- Let the model be the “true” state of the atmosphere, but start with a bad guess (2x the true lightning NO_x emissions)
- We take pseudo-observations of the “true” state at times and locations of IONS06 observations
- Then, let the adjoint assimilate the pseudo-observations and try to converge back to the “true” state from our bad first guess
- Gives confidence that an inversion using these observations will provide a constraint

Ratio of a priori to true LNO_x emissions



... NEXT:

- Perform inversions using real IONS06 observations
- Assess and compare different a priori inventories (Price & Rind, OTD-LIS, WWLLN)

Watch for this in: Ok, this is a work in progress and has no title yet, but it will soon!

Conclusions

(1) Ozone transport to Arctic sites

- Midlatitude sources contribute 25-30% of ozone in the Arctic middle and upper troposphere in the summer
- Adjoint provides complementary detailed source-specific information

(2) Surface ozone in eastern North America

- Adjoint sensitivities can guide further studies into model biases

(3) Midlatitude lightning NO_x emissions constraint

- Simulation experiment shows ozonesonde data can provide a constraint on midlatitude NO_x emissions

Thank you for listening!

EXTRAS

Adjoint Model Methodology

We are concerned with the model misfit to observations, expressed as a quadratic cost function J :

$$J = \frac{1}{2} \langle M(x) - y^{obs}, M(x) - y^{obs} \rangle$$

Performing a linear expansion about some background set of parameters x_o ,

$$J(x) = J(x_o) + \langle \nabla_x J(x_o), x - x_o \rangle$$

$$\delta J = \langle \nabla_x J(x_o), \delta x_o \rangle$$

If M is sufficiently regular, then given a small perturbation δx , the Jacobian $L(x)$ can be used to calculate the perturbation to the state δy :

$$\delta y = L(x_o) \delta x$$

However, using a forward sensitivity study (eg. experimenting with different variations δx) requires an integration of this tangent linear model for each grid point, time interval, iteration $\sim(46 \times 72 \times 55 \times \dots)$

Adjoint Model Methodology

Instead, we define the adjoint model L^* :

$$\langle v, Lw \rangle = \langle L^*v, w \rangle$$

Then, from the derivative of the cost function,

$$\delta J = \langle M(x_o) - y^{obs}, L(x_o)\delta x \rangle$$

$$\delta J = \langle L^*(x_o)(M(x_o) - y^{obs}), \delta x \rangle$$

$$\nabla_x J(x_o) = L^*(x_o)(M(x_o) - y^{obs})$$

The gradient can be used to minimize J by iteratively correcting parameter values. Also, a single calculation of the adjoint model provides sensitivity of model outputs to model inputs:

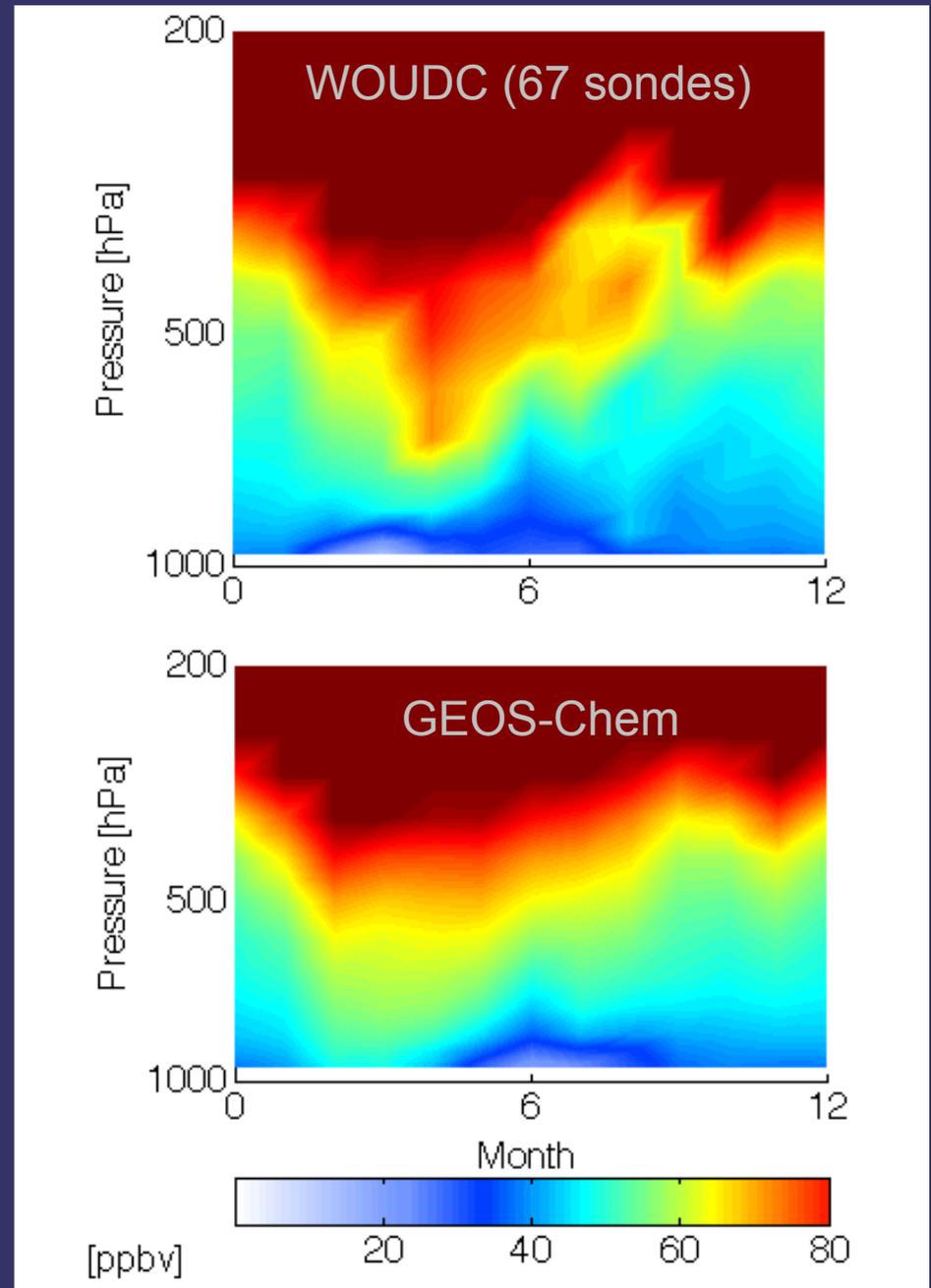
$$\delta x^* = L^*(x_o)\delta y^*$$

Here, we will plot fully normalized sensitivities:

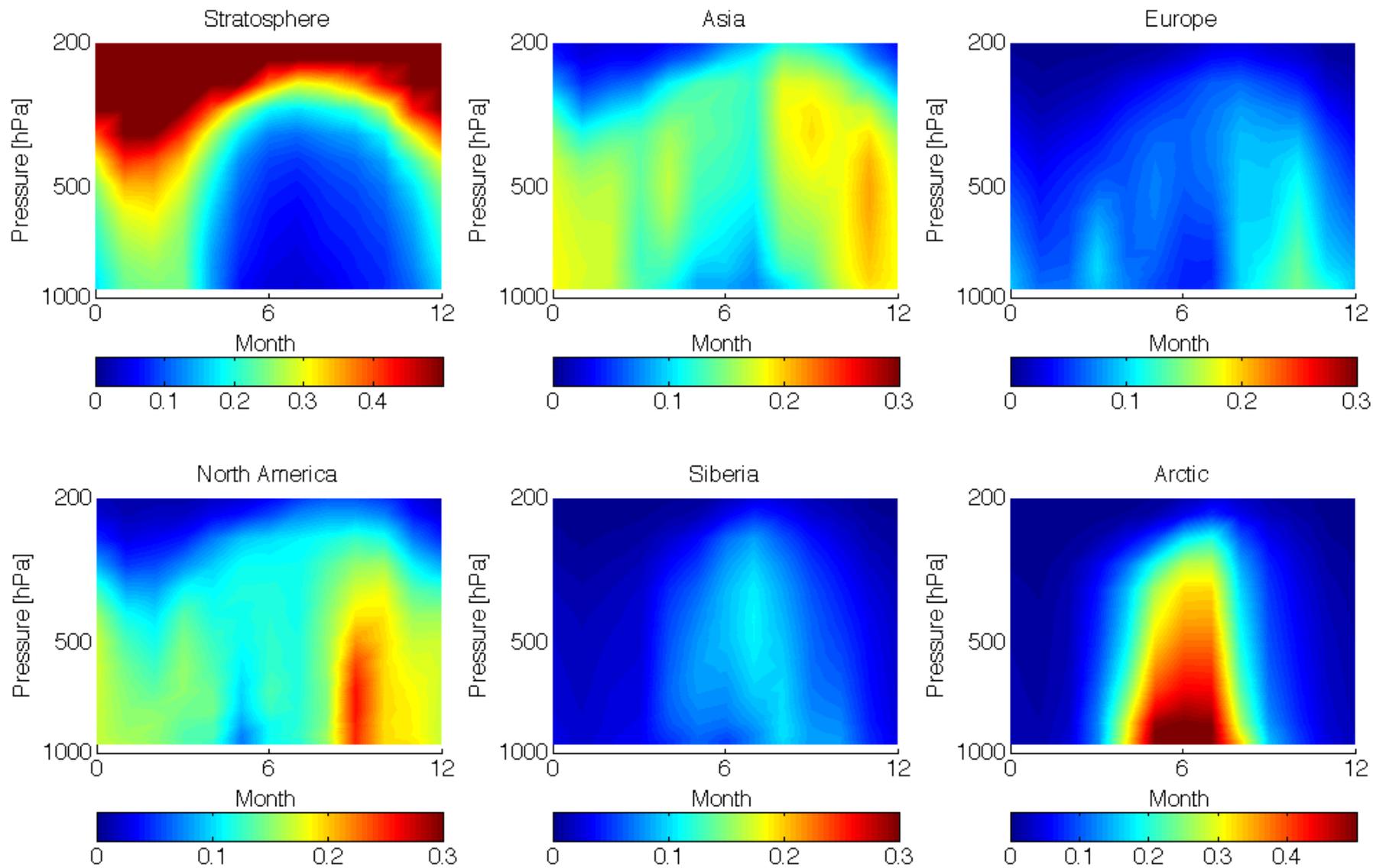
$$\frac{\delta y}{\delta x} \frac{x}{y}$$

Total Ozone Profiles above Eureka, 2006

- model agreement with ozonesondes is good except for a ~10 ppbv underestimate in the upper troposphere in summer
- spring ozone maximum appears earlier in model than in observations

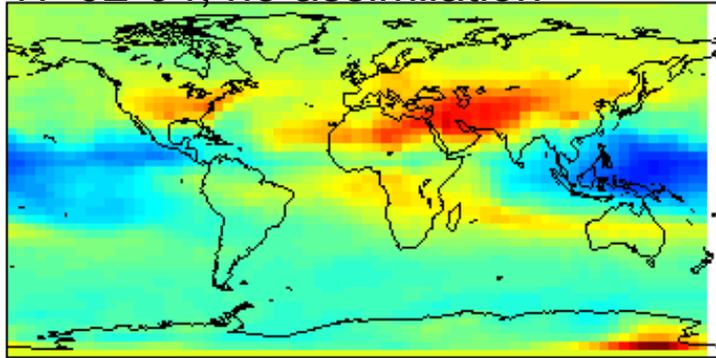


Tagged Ox Profiles above Eureka, 2006

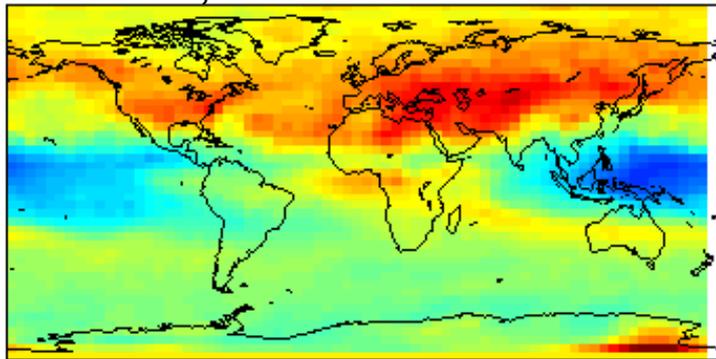


New LNO_x Emissions Produce More Realistic Ozone Distribution

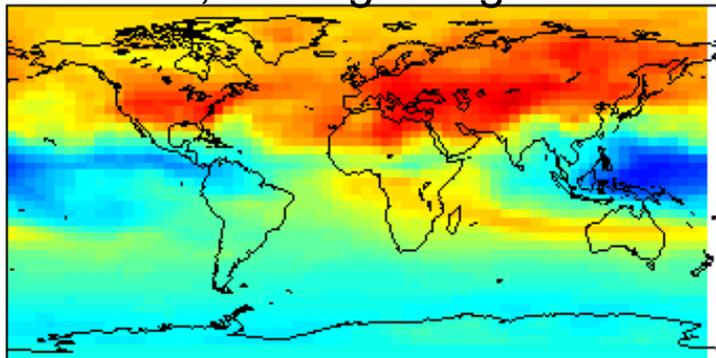
v7-02-04, no assimilation



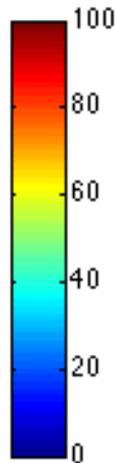
v7-02-04, TES assimilation



v8-01-04, new lightning

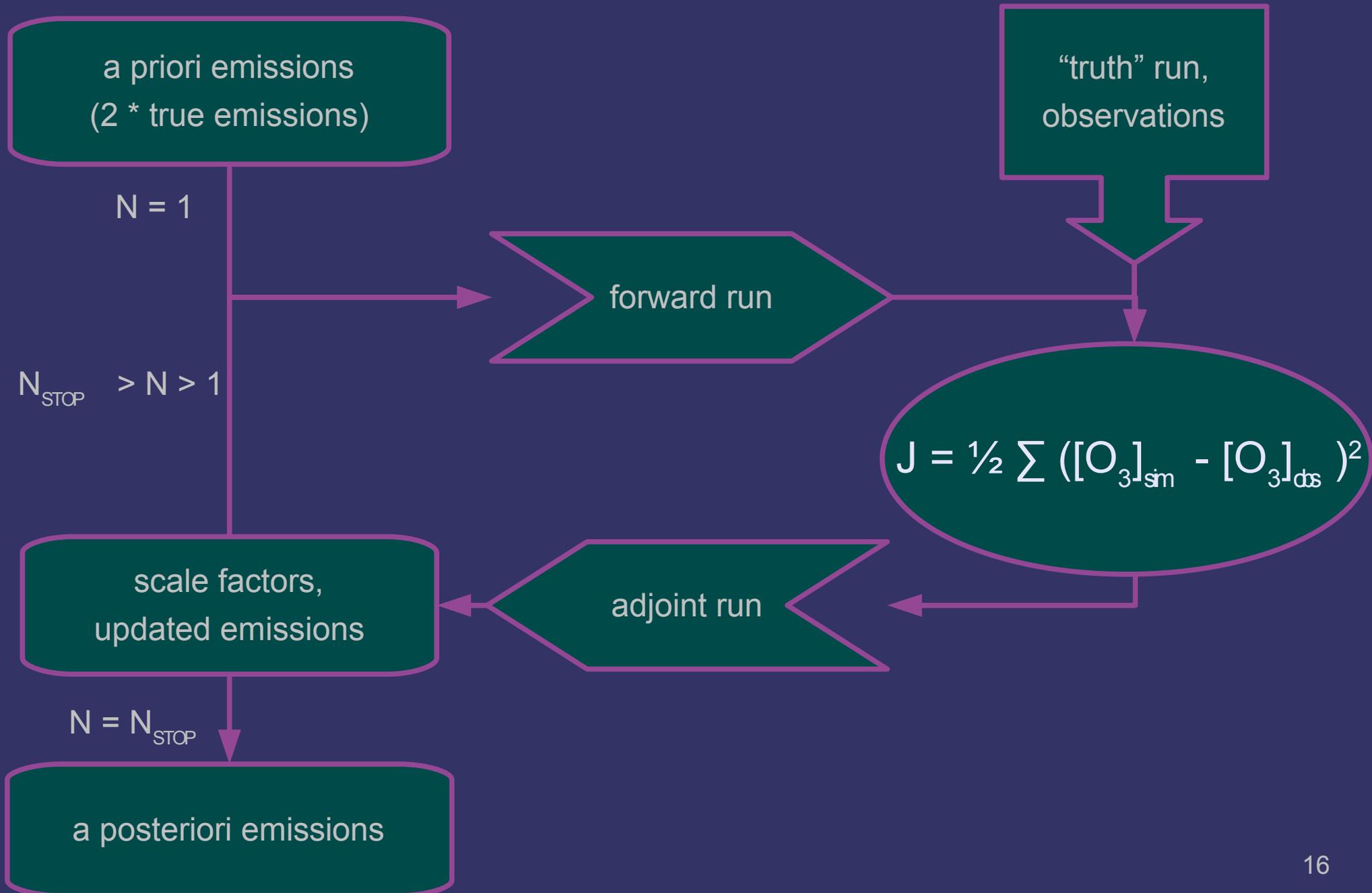


[ppbv]



- Ozone concentrations at 6km
- Relative to N. American ozonesondes, TES assimilation reduces bias from a maximum of -35% to less than 5% between 300-800 hPa
- Relative to the TES assimilation, the new lightning NO_x source reduces the mean bias in the middle troposphere from -7 ppbv to 2 ppbv
- This better representation of midlatitude ozone provides an improved boundary condition for transport into the Arctic

OSSE Evaluation Diagram



References

- Fiore, A. M. et al. (2002) Background ozone over the United States in summer: Origin, trend, and contribution to pollution episodes. JGR, 107 (D15) 10.1029/2001JD000982
- Parrington, M. et al. (2009) Impact of the assimilation of ozone from the Tropospheric Emission Spectrometer on surface ozone across North America. GRL, 36 (L04802) 10.1029/2008GL036935
- Henze, D. K., A. Hakami and J. H. Seinfeld (2007) Development of the adjoint of GEOS-Chem. ACP, 7, 2413-2433
- Price, C. and Rind, D. (1992), A simple lightning parametrization for calculating global lightning distributions, JGR, 97, 9919-9933
- Chai, T. et al. (2006) Chemical data assimilation of Transport and Chemical Evolution over the Pacific (TRACE-P) aircraft measurements. JGR, 111 (D02301) 10.1029/2005JD005883