

Rethinking of Black-Carbon Associated Organic Particles: Insights into Aged Biomass Burning Organic Aerosol

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ABSTRACT

Black carbon (BC) particles in Beijing summer haze play an important role in regional radiation balance and related environmental processes. Understanding the factors that lead to variability in the impacts of BC remains limited. Here, we present observations by a soot-particle aerosol mass spectrometer of BC-containing submicron particulate matter (BC-PM₁) in the summer of 2017 in Beijing, China. These observations were compared to concurrently measured total non-refractory submicron particulate matter (NR-PM₁) by a high-resolution aerosol mass spectrometer (HR-AMS). Distinct variations were observed between NR-PM₁ and BC-PM₁ related to organic aerosol (OA) composition with hydrocarbon-like OA in BC-PM₁ up to two-fold higher than that in NR-PM₁ in fresh vehicle emissions, suggesting that a part of HOA in BC-PM₁ may be undetectable by HR-AMS. Cooking-related OA was only identified in NR-PM₁, whereas aged biomass burning OA (A-BBOA) was a unique factor only identified in BC-PM₁. The A-BBOA was linked to those heavily coated BC, which may lead to enhancement of light absorption ability of BC by a factor of two via the “lensing effect”. More-oxidized oxygenated OA identified in BC-containing particles was found to be slightly different from that observed by HR-AMS, mainly due to the influence of A-BBOA. Overall, these findings highlight that BC in urban Beijing in June is partly of agricultural fire origin and, an unique biomass burning-related OA associated with BC may be ubiquitous in aged BC-PM₁, and this OA may play a role in regional energy balance that has not previously been fully considered.

1 **1. INTRODUCTION**

2 Black carbon (BC) is an important component of atmospheric aerosol that exerts
3 negative effects on regional radiation balance¹ and human health.² It absorbs solar
4 radiation, leading to direct atmospheric heating.³ Indirectly, BC particles can also serve
5 as cloud condensation nuclei upon mixing with hydrophilic species (e.g., sulfate),
6 resulting in changes in cloud properties.⁴ Inhalation of BC is associated with adverse
7 health impacts such as respiratory and cardiovascular diseases, cancers, and birth
8 defects.²

9 Black carbon particles are released to the atmosphere directly from incomplete fuel
10 combustion and open fires.^{1, 3} Non-BC species (e.g., inorganic and organic) can coat
11 onto BC particles in the atmosphere through condensation and/or coagulation
12 processes⁵. These atmospheric processes (also known as “aging”) gradually alter the
13 mixing state and the morphology (e.g., from an externally-mixed fractal structure⁶ into
14 an internally-mixed “core-shell” structure⁷) of submicron BC-containing particles,
15 (BC-PM₁). These alterations can enhance the light absorption capacity of the BC core
16 via the “lensing effect” due to the increased light absorption cross-section as a result of
17 the enhanced coating thickness^{8, 9}. Additionally, the chemical constituents of BC-PM₁
18 may dynamically change during the aging processes, also lead to changes in the light
19 absorption capacity of the particles. Because these physical and chemical processes of
20 both organic and inorganic species inside BC-PM₁ continuously alter particle properties
21 throughout the lifetime of the particles, great uncertainty remains in quantifying the
22 light absorption ability of BC. Understanding the relationship of BC mixing state and
23 chemical composition to the light absorption properties of BC-PM₁, as well as its
24 spatiotemporal distribution, is of importance to accurately evaluate the impacts of BC
25 in regional air quality.

26 Aerodyne high-resolution aerosol mass spectrometry (HR-AMS)¹⁰ has been widely
27 applied in field studies to investigate the chemically-resolved composition of non-
28 refractory submicron particulate matter (NR-PM₁, species that vaporize at temperature
29 < 600 °C)¹¹⁻¹⁷. However, the working temperature of the standard HR-AMS tungsten
30 vaporizer (600 °C) is not sufficient to vaporize refractory species such as BC. To
31 overcome this limitation, soot-particle aerosol mass spectrometry (SP-AMS) is
32 developed¹⁸. In addition to the standard tungsten vaporizer, SP-AMS is equipped with
33 a laser vaporizer (with a wavelength of 1064 nm) which selectively heats BC, together
34 with the non-BC species mixed with it.¹⁹ This novel technique makes it possible to

35 compare the compositions of BC-PM₁ and NR-PM₁, allowing a more accurate
36 assessment of the impacts of BC.

37 There have only been a few studies that have compared the differences of species
38 analyzed by BC-PM₁ and NR-PM₁.^{5, 20, 21} Lee et al.⁵ found that cooking-related organic
39 aerosol (COA) may externally mix with BC in summertime California. The COA factor
40 was identified in NR-PM₁ organic aerosol (OA), but not the BC-related OA. Wang et
41 al.²² found that transported biomass burning organic aerosol could be thickly coated on
42 BC in central Tibetan Plateau and significantly enhance the light absorption capacity of
43 BC cores. Interestingly, the transported biomass burning organic aerosol was not
44 resolved in NR-PM₁ OA particles from concurrent HR-AMS measurements²³. These
45 studies suggest that BC-related OA may undergo different atmospheric processes
46 compared to those do not contain BC.

47 Beijing is a megacity known for high particulate matter (PM) concentrations. BC-
48 PM₁ during haze events of summertime Beijing may have distinct sources and
49 properties than other locations in the world. In this study, we focus on the differences
50 of individual species between BC-PM₁ and NR-PM₁ regarding their chemical
51 composition, mass loadings, sources, and formation pathways in summertime in urban
52 Beijing. Results from this study provide a better understanding of the formation
53 mechanism of OA particles in Beijing haze and valuable insights in assessing their
54 impacts on air quality.

55 **2. EXPERIMENTAL METHODS**

56 **2.1. Sampling site and period**

57 The observations were conducted at a rooftop laboratory (8 m above ground level)
58 in the Tower Division of the Institute of Atmospheric Physics (IAP), Chinese Academy
59 of Sciences (CAS) in urban Beijing (39° 58' N, 116° 22' E), China, from 4 to 29
60 June, 2017. This site has been reported multiple times to be a typical urban observation
61 location²⁴⁻²⁹. The site is located around the North 3rd Ring Road of Beijing. A highway
62 is approximately 360 m to the east and a lot of restaurants (e.g., Sichuan style and BBQ)
63 are within 100 m on the north side.

64

65 **2.2. Instrumentation**

66 Two Aerodyne Aerosol Mass Spectrometers (AMS), including a laser-only Soot-
67 Particle AMS (SP-AMS) and a standard High-Resolution Time of Flight AMS (HR-
68 AMS) were deployed to collect chemical composition and size distributions of BC-PM₁
69 and NR-PM₁, respectively. Three types of species were measured during the campaign:
70 BC-free NR-PM₁ species (Type I), NR-PM₁ species mixed with BC and can be detected
71 by HR-AMS (Type II), and total BC-containing PM₁ including BC cores and all coated
72 species (Type III). HR-AMS is capable of measuring Type I and Type II, while laser-
73 only SP-AMS can measure Type II and Type III. A shared PM_{2.5} cyclone inlet (Model
74 URG-2000-30ED) with 3 Lpm flowrate and a diffusion dryer were used prior to the
75 sampling. The detailed information on the operation of HR-AMS and SP-AMS during
76 the sampling campaign can be found in previous literature^{29, 30}. Details of tuning,
77 calibration, and configurations of the two AMS instruments can be seen in our previous
78 papers^{26, 28, 30}. Mixing ratios of O₃, and NO₂ (Thermo Fisher Scientific, model 49i and
79 model 42C) were measured in parallel simultaneously. Vertical meteorological
80 parameters, including temperature (*T*) and relative humidity (*RH*), were measured from
81 the IAP 325m meteorological tower.

82

83 **2.3. Data Analysis**

84 AMS data analysis was performed by using Squirrel 1.57 and Pika 1.16I based on
85 Igor Pro 6.37 (WaveMetrix Corp.). The measurement of filtered air was performed for
86 24 hours before the start of the campaign to determine the detection limits of various
87 aerosol species and to adjust the fragmentation table. The relative ionization efficiency

88 (RIE) of BC was calibrated with Regal Black (RB, REGAL 400R pigment black, Cabot
89 Corp.). The average ratio of C_1^+ to C_3^+ ionized from pure BC (RB) was determined to
90 be 0.53, which minimizes the influence of C_1^+ from non-refractory organics. The RIE
91 of BC was determined to be 0.17 based on calibrations performed before, in the middle,
92 and at the end of the campaign. RIEs of NO_3^- , SO_4^{2-} , NH_4^+ were determined to be 1.1,
93 0.82, and 3.82, respectively, and default values of 1.3 and 1.4 for RIEs of Chl and Org
94 were applied, respectively.¹⁰ Since direct calibration of NO_3^- , SO_4^{2-} , NH_4^+ are
95 impossible under the laser only configuration, hence, the RIEs calibration of NO_3^- ,
96 SO_4^{2-} , NH_4^+ were all obtained before the tungsten vaporizer was removed by assuming
97 those RIEs were remain unchanged throughout the campaign.²² Polystyrene latex (PSL)
98 spheres (100-700 nm) (Duke Scientific Corp., Palo Alto, CA) were used to calibrate the
99 particle size distribution before the campaign. The default collection efficiency (CE) of
100 0.5 were applied for both HR-AMS and SP-AMS in this study. It should be noticed that,
101 the BC quantification will not be affected by particle bouncing without the tungsten
102 vaporizer, which could affect the CE in the standard HR-AMS measurements.¹⁰
103 However, the CE will be governed by the overlap of particle beam and laser beam.^{5, 21}
104 Both HR-AMS and SP-AMS chemically resolved mass concentrations of NR- PM_1 and
105 BC were calculated based on V-mode high-resolution fitting. Due to different
106 vaporization schemes between the HR-AMS and SP-AMS, mass spectra from these two
107 instruments even for the same population of aerosols are not entirely the same. Because
108 laser-only SP-AMS generally results in less overall fragmentation, its mass profile may
109 contain more large m/z fragments and less small m/z fragments compared to that from
110 HR-AMS²¹. In addition, the elemental ratios of organics reported here, i.e., oxygen-to-
111 carbon and hydrogen-to-carbon ratios (O/C and H/C) were only calculated based the
112 “Improved-Ambient (I-A)” method³¹.

113 Positive matrix factorization (PMF)³² was performed on the high-resolution
114 organic mass spectra matrix of both NR- PM_1 and BC- PM_1 (e.g., BC (C_x^+), NR- PM_1
115 associated with BC) across m/z 12–120 using Evaluation Tool written in Igor³³,
116 following the standard procedure³⁴. Four types of organic aerosol (OA) from total NR-
117 PM_1 (see our previous paper)¹⁶ and five OA factors associated with BC were identified.
118 C_x^+ was involved in the calculation of elemental ratios (e.g, O/C and H/C) of PMF OA
119 factors. All data presented in this paper were averaged hourly and are presented at local
120 time (Beijing Time, UTC+8).

121 3. RESULTS AND DISCUSSION

122 3.1. Overview of observations

123 Figure 1 shows the temporal variations of selected chemical species during the
124 campaign. Information for other variables is provided in the supplementary materials
125 (SM). The two cases labeled in Figure 1 are of particular interest. Case I (June 8-13)
126 was characterized with high NO₂ concentrations (average 26.7 ± 13.5 ppb, Table S1)
127 and relatively low O₃ concentrations (41.7 ± 30.0 ppb) with NO₂-to-O₃ ratio of 0.64.
128 Case II (June 17-22) was featured by low NO₂ (14.9 ± 5.9 ppb) and high O₃ ($84.6 \pm$
129 30.6 ppb) concentrations with an NO₂-to-O₃ ratio of 0.18. Unlike winter Beijing haze
130 pollution, *RH* remained at a relatively low level ($36.5 \pm 15.3\%$), which is not expected
131 to play a significant role in OA formation during the campaign (Figure 1b and Figure
132 S1). In contrast, a strong correlation has been observed between temperature and O₃ (r^2
133 = 0.53). The temperature was higher on average in Case II (29.8 ± 3.8 °C) than in Case
134 I (26.1 ± 4.1 °C).

135 The mass concentrations and mass concentration ratios of organic (Org), sulfate
136 (SO₄²⁻) and nitrate (NO₃⁻) in NR-PM₁ (in solid line) and BC-PM₁ (in dotted line) are
137 shown in Figures 1c-e. High correlations were observed between BC-PM₁ and NR-PM₁
138 measurements for SO₄²⁻ ($r^2 = 0.70$) and NO₃⁻ ($r^2 = 0.86$), but not for Org ($r^2 = 0.49$).
139 This result suggests that, BC-PM₁ Org has distinct sources or formation pathways from
140 NR-PM₁ Org. Comparing two cases, the average mass ratios of BC-PM₁ to NR-PM₁ for
141 SO₄²⁻ and NO₃⁻ in Case I (0.24 ± 0.11 and 0.37 ± 0.12) were close to those in Case II
142 (0.19 ± 0.06 and 0.31 ± 0.07). However, ratios of BC-PM₁ to NR-PM₁ for Org were
143 160% greater for Case I (0.74 ± 0.32) compare to Case II (0.46 ± 0.13). During the
144 nighttime, this ratio increases to almost unity in Case I. Additionally, BC concentration
145 in Case I (average 2.6 ± 1.6 μg m⁻³) was 1.5 folds higher than in Case II (average $1.7 \pm$
146 0.8 μg m⁻³). The implication is that the organic is mostly associated with BC and likely
147 comprised of freshly emitted compounds in Case I. This is also evident by the moderate
148 correlation between NO₂ and BC-PM₁ Org ($r^2 = 0.42$) in Case I. On the other hand, the
149 lower Org ratio in Case II with higher O₃ concentrations indicates greater oxidation and
150 secondary processes in non-BC particles.

151 3.2. Source apportionment of BC-PM₁ OA

152 To further investigate the differences between organics in NR-PM₁ and BC-PM₁,
153 the comparison of PMF OA factors between NR-PM₁ and BC-PM₁ Org is necessary.
154 Four factors were identified from PMF analysis of the NR-PM₁ Org matrix, including
155 hydrocarbon-related OA (HOA), cooking OA (COA), less-oxidized oxygenated OA
156 (LO-OOA), and more-oxidized oxygenated OA (MO-OOA). Details of the NR-PM₁
157 PMF analysis can be found in our previous paper³⁰. Here we only present the PMF
158 results of the SP-AMS measured BC-PM₁ Org. As shown in Figure 2, five OA factors
159 were resolved by PMF with factors including a HOA, a less oxidized OOA (OOA1 or
160 LO-OOA), three more-oxidized OOA (MO-OOA). The MO-OOA factor includes an
161 aged BBOA (A-BBOA), a sulfate- and photochemical-related OOA (OOA2), and a
162 nitrate-related OOA (OOA3). A key diagnostic plot of this PMF solution is presented
163 in Figure S2.

164 BC-PM₁ HOA consists of a series of hydrocarbon fragments (C_xH_y⁺) in its mass
165 spectrum (Figure 2f), thus having a low O/C ratio (0.13) but high H/C ratio (1.62). It
166 has a *r*² of 0.92 with C₄H₉⁺ (m/z = 57) and a *r*² of 0.57 with NO_x (Figure 2a), indicative
167 of its sources from vehicle emissions³⁵. It also correlated tightly with BC (*r*² of 0.70)
168 and a series of polycyclic aromatic hydrocarbons (PAHs) ions, e.g., C₉H₇⁺ (m/z 115, *r*²
169 of 0.63).

170 The second factor has a remarkably high fraction of the BBOA marker ions of
171 C₂H₄O₂⁺ (m/z = 60) (1.31%) and C₃H₅O₂⁺ (m/z = 73) (1.34%) in its mass spectrum
172 (Figure 2g), much higher than that observed in non-BBOA (e.g., 0.3% at m/z = 60) in
173 previous studies^{22, 35, 36}. As expected, the temporal variation of this factor correlated
174 tightly with those of C₂H₄O₂⁺ and C₃H₅O₂⁺ (*r*² of 0.71 and 0.72, respectively). In
175 addition, the mass spectrum of this factor is strikingly similar to that of the transported
176 BBOA which was observed at a remote site in the central Tibetan Plateau²², with a *r*²
177 of 0.97. Here we categorized the transported BBOA as A-BBOA identified in this study.
178 Similar to the A-BBOA observed in Tibetan Plateau, which has an O/C ratio of 0.51,
179 this factor also has a relatively high O/C ratio of 0.48, greater than that of primary
180 BBOA (O/C of 0.18–0.26)²². These findings support that the second factor may be
181 associated with the oxidation of biomass burning emissions. The temporal variation of
182 ABBOA in the Tibetan Plateau was reported to be highly correlated with the potassium
183 ion fraction (K⁺, *r*² of 0.78), and K₃SO₄⁺ (*r*² of 0.92). However, the temporal variation
184 of the second factor in this study is only correlated well with that of K₃SO₄⁺ (*r*² of 0.64)

185 but not K^+ (r^2 of 0.01). The reason for this phenomenon is that the major source of K^+
186 in remote sites like the Tibetan Plateau was long-distance transport of K_2SO_4 particles,
187 which probably from biomass burning-related K-containing salts interacts with H_2SO_4 .
188 In contrast, there are multiple primary sources of K^+ in PM_{10} (e.g., diesel-vehicle
189 emissions, and mainly KCl particles) in urban areas (Figure S3). Based on these
190 observations, $K_3SO_4^+$ could be defined as a A-BBOA marker ion. Hence, this second
191 factor is identified as A-BBOA that was subjected to oxidation during transport to the
192 measurement area as presented in the bivariate polar plot and three-day back trajectories
193 (Figure S4). June should be the month of maximum agricultural-related biomass
194 burning in the North China Plain, although we thought that this burning had been
195 banned in recent years because of air quality concerns.³⁷ The implication is that the
196 effectiveness of banning straw burning may be overestimated.

197 The OOA1 factor has an O/C of 0.28 (Figure 2h). Similar to the NR- PM_{10} LO-
198 OOA¹⁶, it is highly correlated with $C_2H_3O^+$ (r^2 of 0.72). The $C_2H_3O^+$ ion ($m/z = 43$) is
199 an important component of secondary organic aerosol (SOA)^{20, 38} and the diurnal
200 patterns of the OOA1 and $C_2H_3O^+$ both show a great enhancement around noontime
201 (Figure S5), indicating the importance of secondary formation of less oxidized organic
202 aerosol through daytime photochemical activity.

203 The OOA2 factor has an O/C of 0.42 (Figure 2i) and the OOA3 factor has a smaller
204 O/C of 0.32 (Figure 2j). OOA2 correlated strongly with sulfate (r^2 of 0.92; Figure 2d)
205 and OOA3 correlated highly with nitrate (r^2 of 0.97; Figure 2e). These features agree
206 well with the previously observation for low-volatility OOA (sulfate-related OOA) and
207 semi-volatile OOA (nitrate-related OOA) in Tibetan Plateau²².

208 3.3. Comparison of NR-PM₁ and BC-PM₁ OA factors

209 The sum of the above-mentioned BC-PM₁ A-BBOA, OOA2, and OOA3 fractions
210 is comparable to the NR-PM₁ MO-OOA factor, based on their high O/C ratios. Figures
211 3a-c are comparisons of the mass loadings of HOA, LO-OOA, and MO-OOA in both
212 NR-PM₁ and BC-PM₁. NR-PM₁ HOA, LO-OOA, and MO-OOA are strongly correlated
213 with their counterpart fractions of BC-PM₁, with r^2 values of 0.68, 0.60, and 0.61,
214 respectively. In Case I, most of the time, the mass loadings of BC-PM₁ HOA and MO-
215 OOA are higher than those in NR-PM₁, while LO-OOA shows the opposite trend. In
216 Case II, the mass loadings of BC-PM₁ HOA are also generally higher than those of NR-
217 PM₁ HOA, however, NR-PM₁ MO-OOA and LO-OOA are almost two folds higher than
218 those of BC-PM₁. Figures 3d-f are comparisons of the fractions of HOA, LO-OOA, and
219 MO-OOA in NR-PM₁ and non-BC material in BC-PM₁ (coatings), respectively. In Case
220 I, the fractions of HOA and MO-OOA internally-mixed with BC are almost two times
221 and four times higher, respectively, than those in NR-PM₁, whereas the two LO-OOA
222 fractions closely track each other. In Case II, two LO-OOA fractions are still overlapped,
223 but compared to Case I, the fraction of HOA in BC-PM₁ coatings is over four times that
224 of NR-PM₁ HOA, and the difference between the two MO-OOA fractions is smaller.

225 As shown in Figure 4, the average of BC-PM₁ HOA fractions (0.27 ± 0.17 and 0.11
226 ± 0.07 , respectively) are higher than those in NR-PM₁ (0.12 ± 0.08 and 0.02 ± 0.02 ,
227 respectively) in both Case I and Case II, indicating that HOA particles is more internally
228 mixed with BC compared to other OA materials. There is a higher fraction of low-
229 volatility HOA in BC-PM₁ that can only be evaporated under higher temperature
230 conditions (> 600 °C) in SP-AMS. It is also possible that some species (e.g., PAHs)
231 could be buried inside of the BC coatings or trapped in nano-carbon spheres³⁹, which
232 are not readily evaporated by the tungsten vaporizer in HR-AMS. However, the
233 possibility that RIE of OA coating may be lower than the default RIE value should also
234 be considered. Nevertheless, previously reported HOA levels by HR-AMS need to be
235 reinvestigated due to the possible underestimation of the refractory of HOA.

236 The average mass loadings of NR-PM₁ LO-OOA in both Case I and Case II were
237 higher than those of BC-PM₁. However, the fraction of LO-OOA in both NR-PM₁ and
238 BC-PM₁ coatings were very close to each other during the two cases, with an average
239 value of 0.23 ± 0.10 and 0.25 ± 0.12 , respectively, indicating that the probability of LO-
240 OOA condensation onto the two different types of particles is similar.

241 A greater difference between the MO-OOA fractions in NR-PM₁ and BC-PM₁ was

242 observed in Case I than in Case II, and there is more MO-OOA in BC-PM₁ than in NR-
243 PM₁ in Case I. A similar comparison between NR-PM₁ MO-OOA with BC-PM₁ MO-
244 OOA without A-BBOA can be found in SI (Figure S6), which shows closer fractions in
245 both Case I and Case II. Therefore, one possibility which may cause higher MO-OOA
246 fraction in BC-PM₁ than that in NR-PM₁ in Case I is the presence of the BC-PM₁ A-
247 BBOA, which is only identified from the BC-PM₁ OA. More details of the BC-PM₁ A-
248 BBOA are discussed in Section 3.4.

249 3.4. Characteristics of A-BBOA in BC-containing PM₁

250 Figure 5 shows the high-resolution mass spectra of A-BBOA observed in Nam Co
251 (June 2015) and Beijing (June 2017) by laser-only SP-AMS. A mass spectra very
252 similar to that observed in Beijing was also observed in Nanjing (February 2017)⁴, with
253 a r^2 of 0.95. The A-BBOA observed in Nam Co (the Tibetan Plateau) was found in the
254 thickest coated and internally-mixed BC-PM₁ (the mass ratio of coatings to BC core
255 (R_{BC}) can reach 14), which enhances the light absorption ability (E_{abs}) of the BC core
256 by a factor of 1.5 to 2.0 via the “lensing effect”.

257 As shown in Figure 6, A-BBOA was associated with those large particles ($D_{va} >$
258 300nm) which were also heavily-coated ($R_{BC} > 9$, Figure 6a and 6c). Because A-BBOA
259 is a moderately aged OA, the OSc was very steady when $R_{BC} > 9$ (Figure 6c). Figure
260 6b presents the fractions of the OA factors (left) and the degree of light absorption
261 enhancement (E_{abs} , estimated by the mass ratios of BC measured by aethalometer model
262 33 and SP-AMS), as a function of R_{BC} . Figure 6d shows the temporal variations of the
263 fractions of NR-PM₁ OA and BC-PM₁ OA from 15:00 to 24:00 on June 17, 2017 when
264 the highest A-BBOA concentrations were observed. There is a significant enhancement
265 of A-BBOA which may account for up to 60% of the total OA coatings, which could
266 enhance the BC-PM₁ MO-OOA fraction (within the purple frame in the bottom panel
267 of Figure 6d).

268 In this study, A-BBOA was only observed by SP-AMS and was indeed only
269 associated with BC. It is likely that A-BBOA was emitted together with BC when
270 burning biomass fuel, and was oxidized subsequently during the transport. However,
271 we cannot exclude the possibility that A-BBOA can be detected by HR-AMS. For
272 example, it might be included in NR-PM₁ MO-OOA factor. Without separating A-
273 BBOA from other organic species, the source apportionment for HR-AMS may obscure
274 air-quality- and climate-related implications of A-BBOA in the atmosphere, such as the
275 enhancement of aerosol light absorption ability (Figure 6b).

276 **4. ATMOSPHERIC IMPLICATIONS**

277 Online chemical characteristics of BC and its associated species was for the first
278 time elucidated in urban Beijing in summer, and compared with those of NR-PM₁
279 species. The biggest difference between the two measurements was in the composition
280 of the organic species. In particular, we found BC in urban Beijing in June is partly of
281 agricultural fire origin and, an unique biomass burning-related OA factor (A-BBOA)
282 which was moderated aged, only existed in thickly coated BC-PM₁ ($R_{BC} > 9$), but not
283 NR-PM₁. The unique A-BBOA could make up a significant portion of BC coating
284 material. In addition to Beijing, the A-BBOA was also identified in other locations (such
285 as central Tibet Plateau²² and Nanjing⁴), suggesting that it may be ubiquitously presence
286 in BC-PM₁ in the atmosphere.

287 BBOA species are known to constitute a large portion of light-absorbing organics
288 (brown carbon, BrC). The delay of BBOA oxidation and its longer duration time on BC
289 cores can extend the impacts of BC. Moreover, together with our previous study of BC-
290 associated A-BBOA in Tibet, results presented herein demonstrate that A-BBOA could
291 lead to thick coating on BC cores, meaning a significant “lensing effect” to the
292 enhancement of BC light absorption⁴⁰. As a key component of BC coating, presence of
293 this factor may also alter the bulk hygroscopicity of BC-PM₁. It could therefore affect
294 its activation ability as cloud condensation nuclei (CCN)⁴. Overall, the emission,
295 evolution and transport of such A-BBOA, may significantly affect the atmospheric
296 behaviors and influence the role of BC in the air quality and climate (e.g., radiative
297 forcing and precipitation). Future laboratory, field, and modeling studies are required
298 to evaluate the regional environmental impacts of A-BBOA.

ASSOCIATED CONTENT

Supporting Information

AUTHOR INFORMATION

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ABBREVIATIONS

BC Black carbon

PM₁ Particulate matter with an aerodynamic diameter smaller than 1 μm

NR-PM₁ non-refractory PM₁

BC-PM₁ BC-containing particles in PM₁

BrC Brown carbon

HR-AMS High-resolution aerosol mass spectrometer (Aerodyne Research Inc.)

SP-AMS Soot-particle aerosol mass spectrometer (Aerodyne Research Inc.)

IE Ionization efficiency

RIE Relative ionization efficiency

HRMS High-resolution mass spectra

PMF Positive matrix factorization

OA Organic aerosol

SOA Secondary organic aerosol O/C Oxygen-to-carbon ratio

H/C Hydrogen-to-carbon ratio

A-BBOA Aged biomass burning organic aerosol

SV-OOA Semi-volatile oxygenated organic aerosol

LV-OOA low-volatility oxygenated organic aerosol

MO-OOA more-oxidized oxygenated organic aerosol

LO-OOA less-oxidized oxygenated organic aerosol

R_{BC} mass ratio of BC coatings to BC

D_{va} Vacuum aerodynamic diameter

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List of Figures

TOC

Figure 1. Temporal variations of selected chemical species measured in Beijing on June 4 -29, 2017. (a) mixing ratios of nitrogen dioxide (NO_2) and ozone (O_3); (b) 15-m relative humidity (RH) and temperature (T); (c-e) on the left are the mass loadings of organic (Org), sulfate (SO_4^{2-}) and nitrate (NO_3^-) measured by HR-AMS and SP-AMS, and on the right are mass ratios of individual BC- PM_{10} species to NR- PM_{10} species (e.g., BC- PM_{10} Org to NR- PM_{10} Org). The NR- PM_{10} species measured by HR-AMS is in solid line, and the BC- PM_{10} species measured by SP-AMS is in the dotted line. The shaded areas are raining periods. The observation period is divided into two cases according to the mixing ratio of nitrogen NO_2 , Case I and Case II, which represent high NO_2 and low NO_2 mixing ratios, respectively.

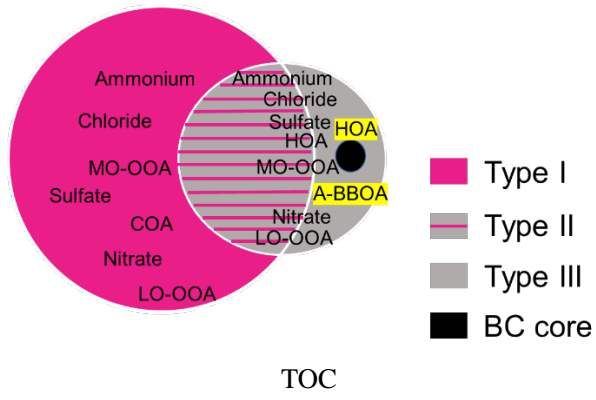
Figure 2. Temporal variations (left panels), high-resolution mass spectra (right panels) of five OA factors in summer 2017: (a) and (f) HOA, (b) and (g) A-BBOA, (c) and (h) OOA1 (LO-OOA), (d) and (i) OOA2, and (e) and (j) OOA3. Also shown in the left panels are the time series of other tracers, including C_4H_9^+ , NO_x , $\text{C}_2\text{H}_4\text{O}_2^+$, K_3SO_4^+ , $\text{C}_6\text{H}_{10}\text{O}^+$, $\text{C}_2\text{H}_3\text{O}^+$, SO_4^{2-} and NO_3^- .

Figure 3. Temporal variations of NR- PM_{10} and BC- PM_{10} (a-c) HOA, LO-OOA, and MO-OOA (left panels) and (d-e) their fractions. NR- PM_{10} OA factors are in red, and the BC- PM_{10} OA factors are in black. Here BC- PM_{10} MO-OOA is the sum of A-BBOA, OOA2 (sulfate-related OOA), and OOA3 (nitrate-related OOA).

Figure 4. Box plots of mass loadings and fractions of five selected species (HOA, LO-OOA, MO-OOA, SO_4^{2-} , and NO_3^-) in Case I and Case II. The bounds of boxes represent quartiles, the whiskers indicate the 90th and 10th percentiles, and the lines and dots inside the boxes are median and mean values. NR-PM₁ OA factors are in red, and the BC-PM₁ OA factors are in black.

Figure 5 Comparison between the high-resolution mass spectra of A-BBOA obtained in Nam Co (June 2015) and Beijing (June 2017).

Figure 6. (a-c) the mass loadings of BC, BC-PM₁ Org, fractions of BC-PM₁ OA factors, E_{abs} , the oxidation state ($\text{OSc} = 2*(\text{O}/\text{C}) - (\text{H}/\text{C})$) of BC-PM₁ Org, and the size distribution of BC-PM₁ Org as a function of coating thickness (R_{BC}). (d) temporal variations of OA fractions of NR-PM₁ and BC-PM₁ from 15:00 to 24:00 on June 17, 2017.



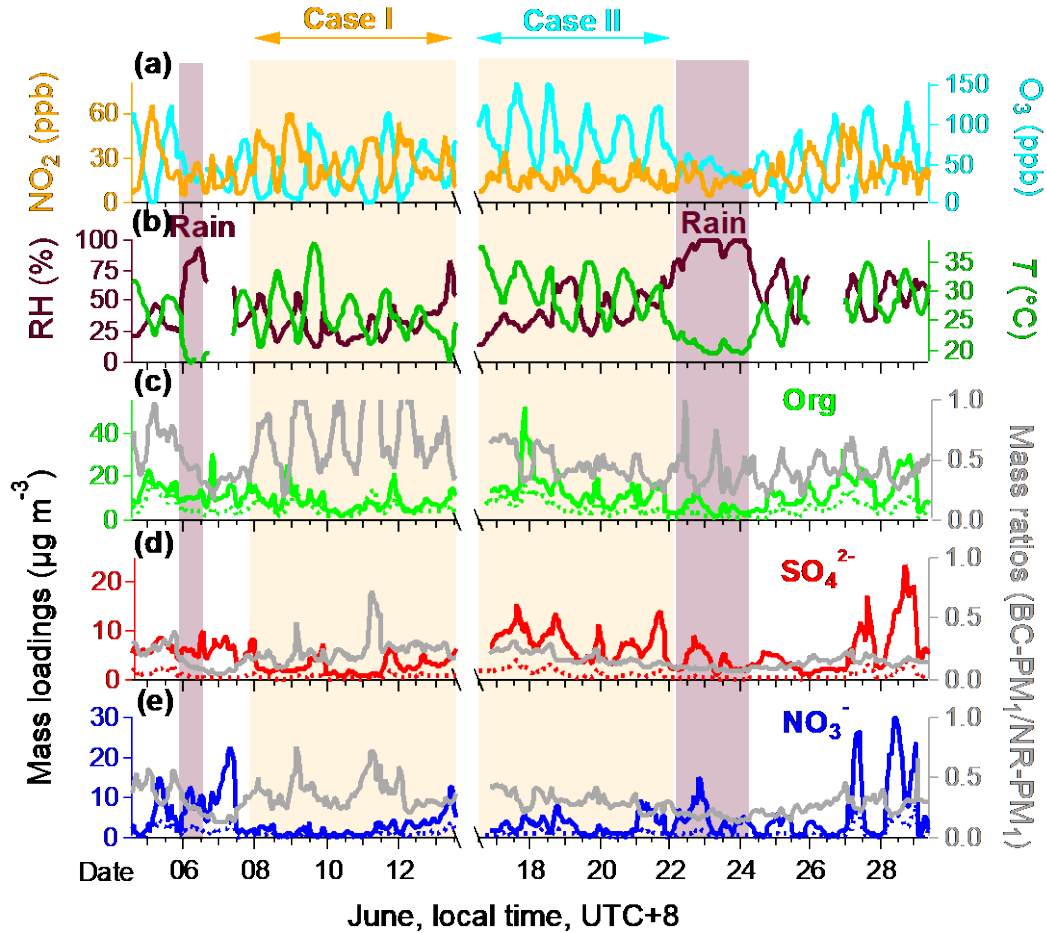


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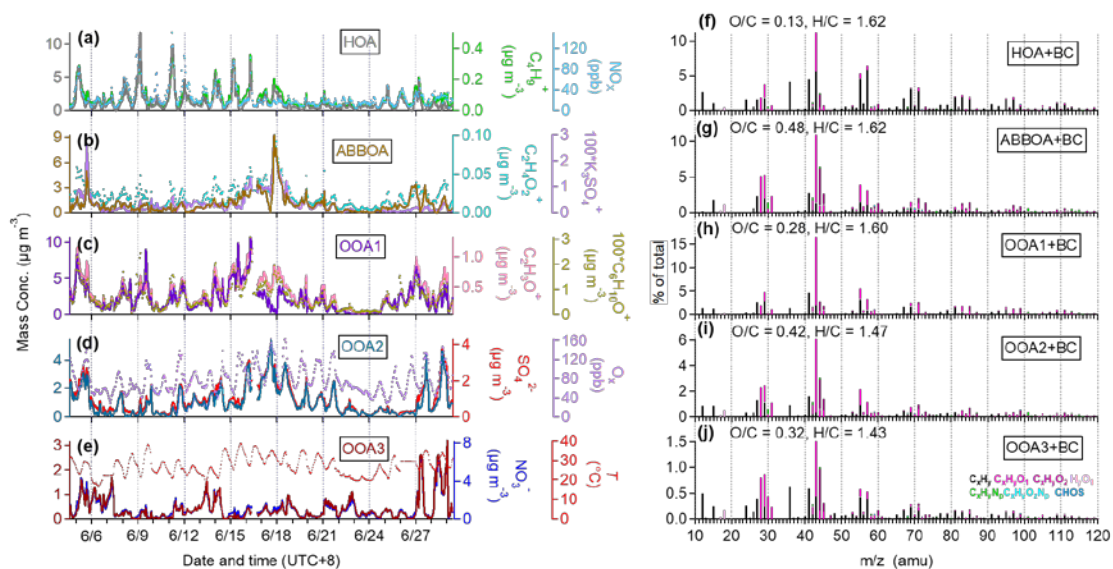


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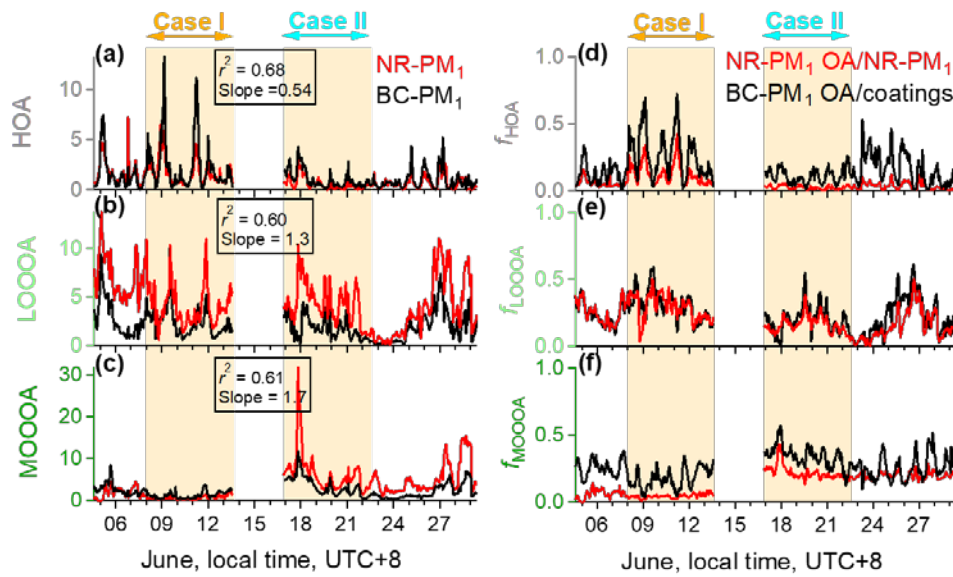


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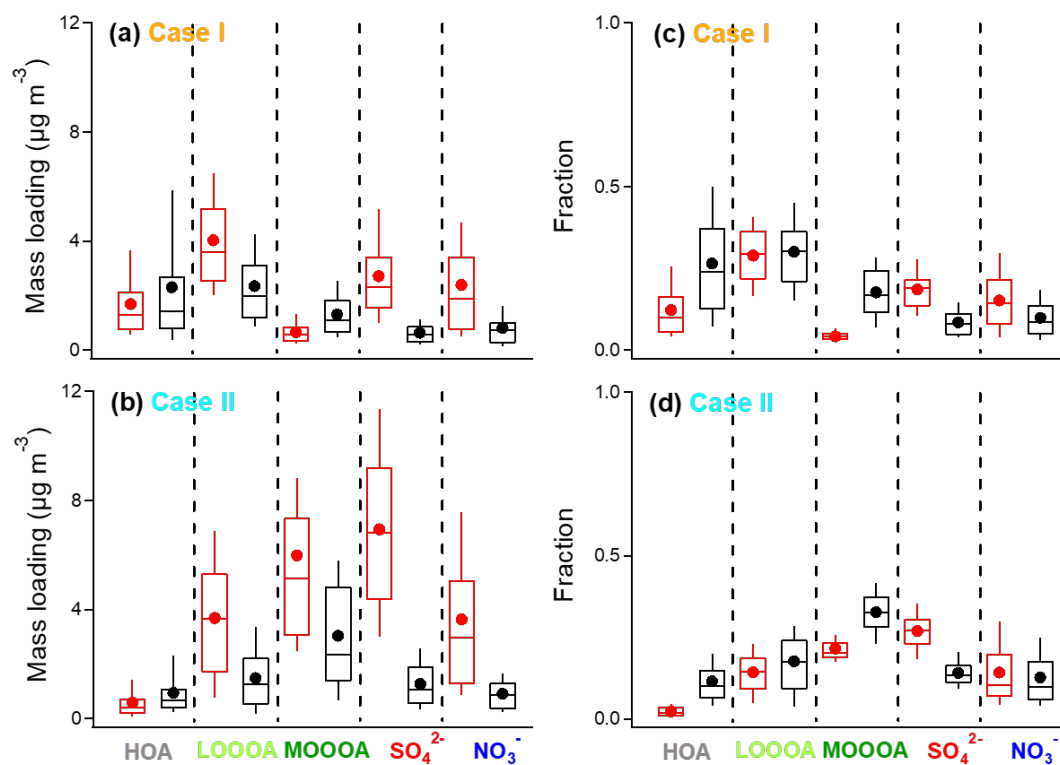


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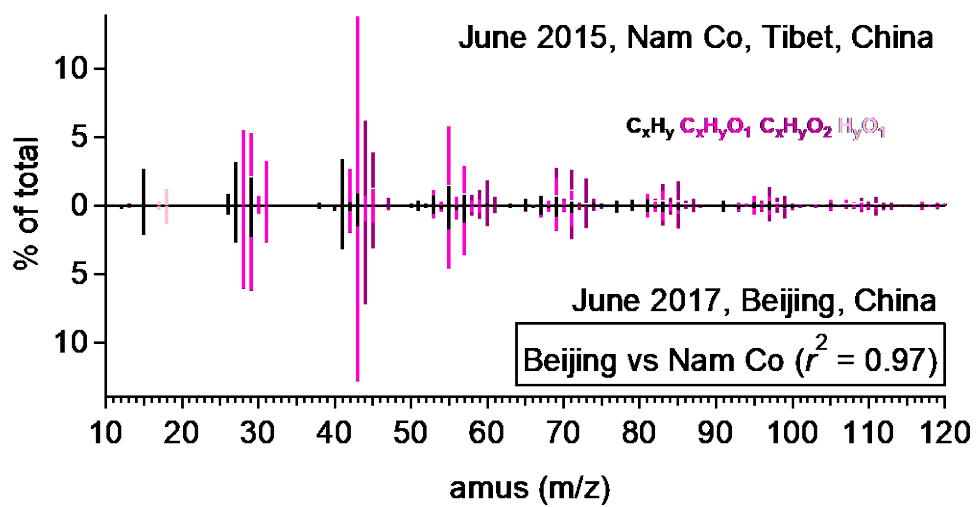


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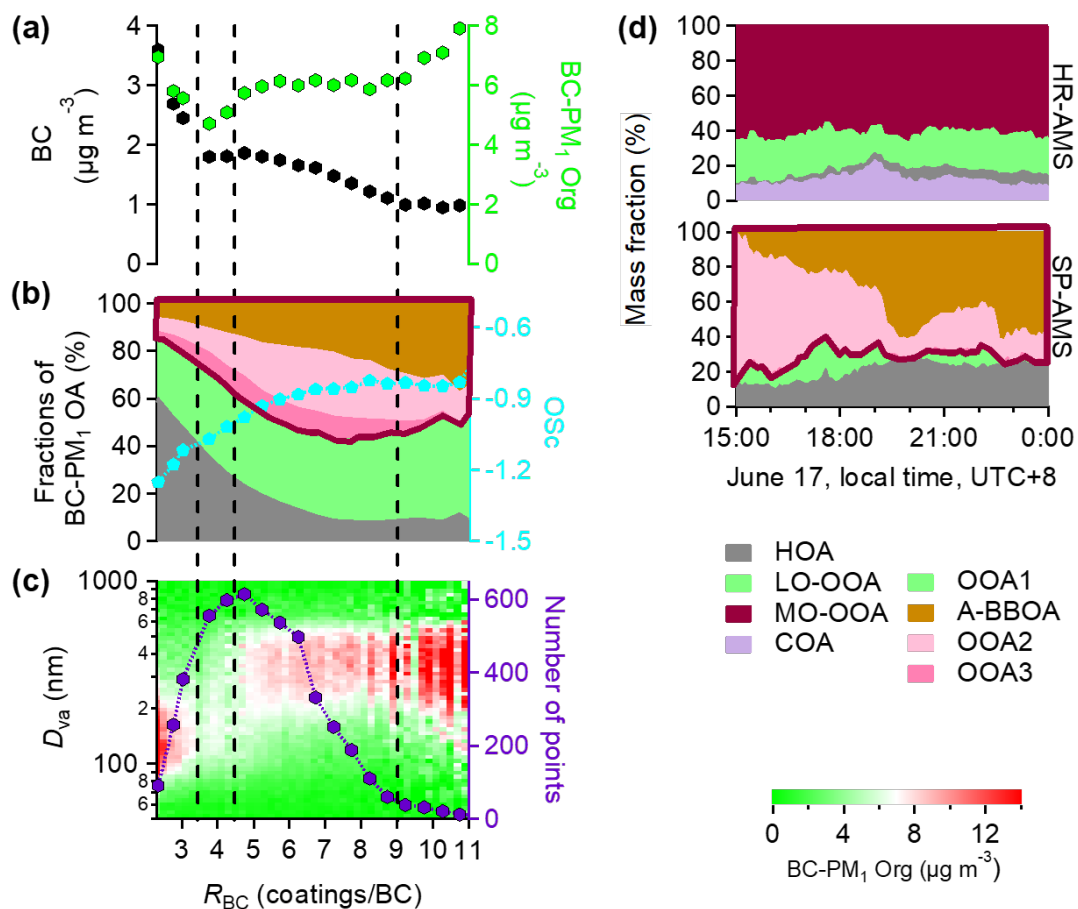


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